

**Systematics of run-up distributions from
dislocation and landslide sources:
A near-field discriminant**

Emile A. Okal

Department of Geological Sciences

Northwestern University

Evanston, IL 60201

Costas E. Synolakis

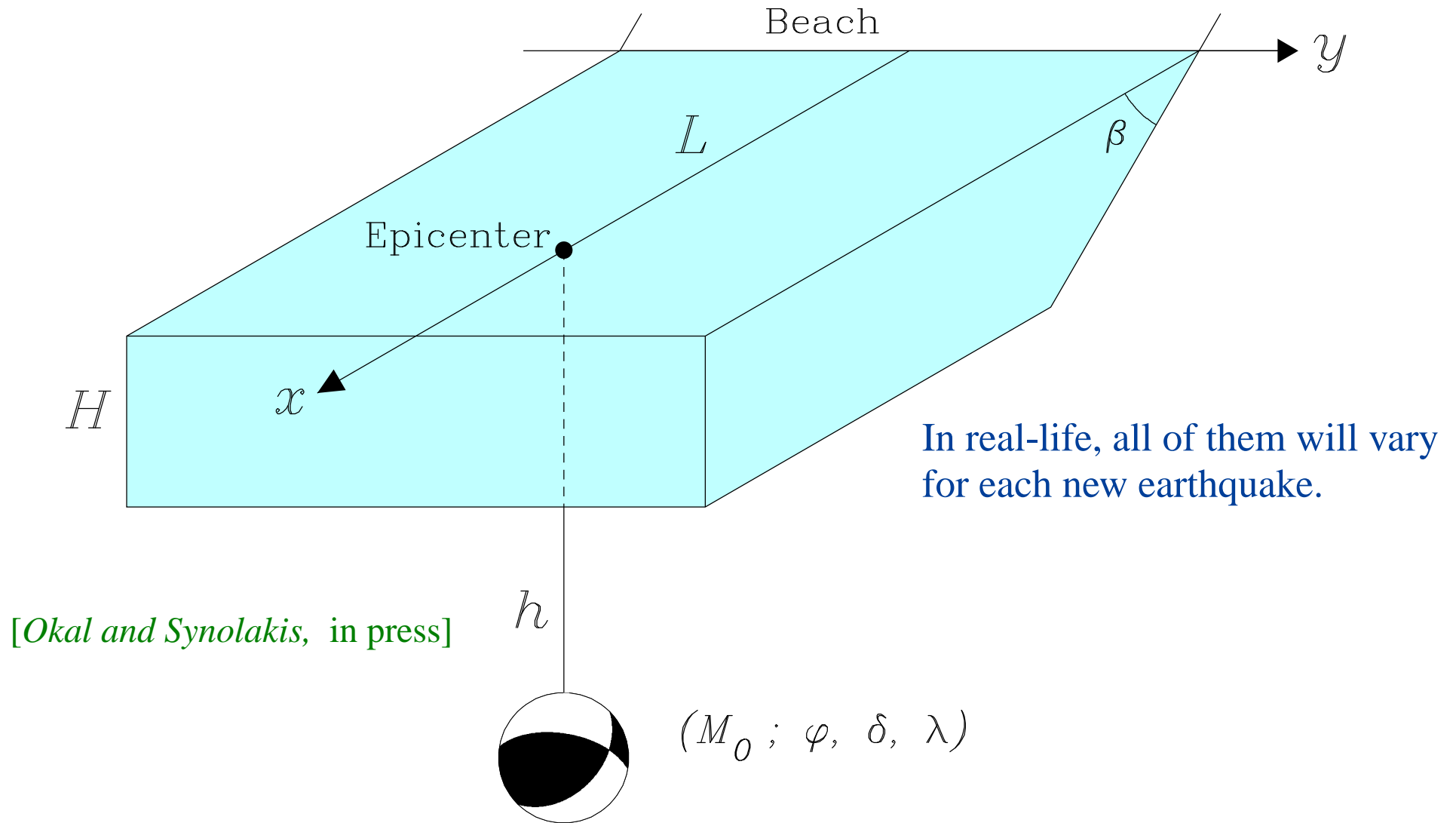
Department of Civil Engineering

University of Southern California

Los Angeles, CA 90089

THE DISLOCATION SOURCE in the NEAR FIELD

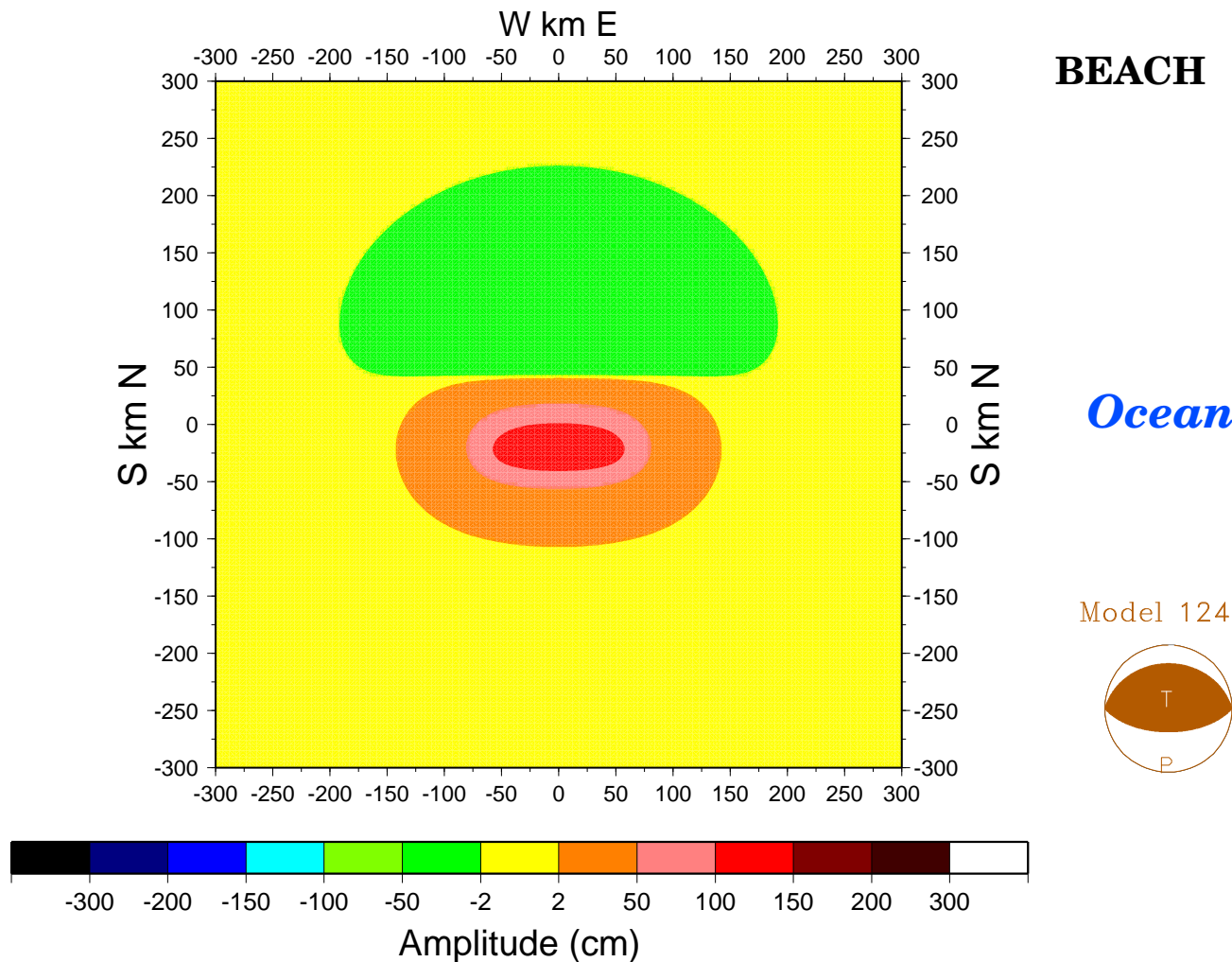
A full description requires at least 8 parameters.



We explore systematically their influence on run-up and seek to define INVARIANTS

NEAR-FIELD: *The Earthquake Dislocation*

- Compute Ocean-Bottom Deformation due to Dislocation



- Simulate Tsunami Propagation to Beach and Run-up

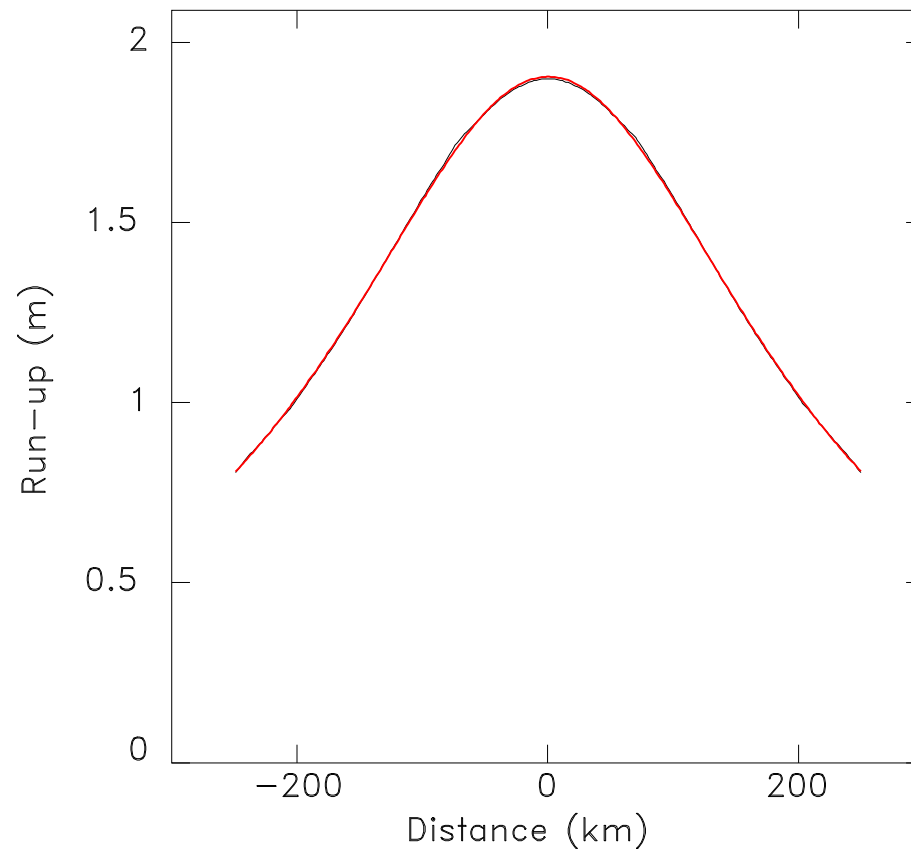
- Simulate Tsunami Propagation to Beach and Run-up

- Fit Bell Curve

$$\zeta = \frac{b}{\left(\frac{x-c}{a}\right)^2 + 1}$$

- Retain aspect ratio $I = b/a$

Aspect ratio $I_2 = b/a = 0.89\text{E-}05$



$a = 214.2 \text{ km}$

$b = 1.90 \text{ m}$

$c = 0.4 \text{ km}$

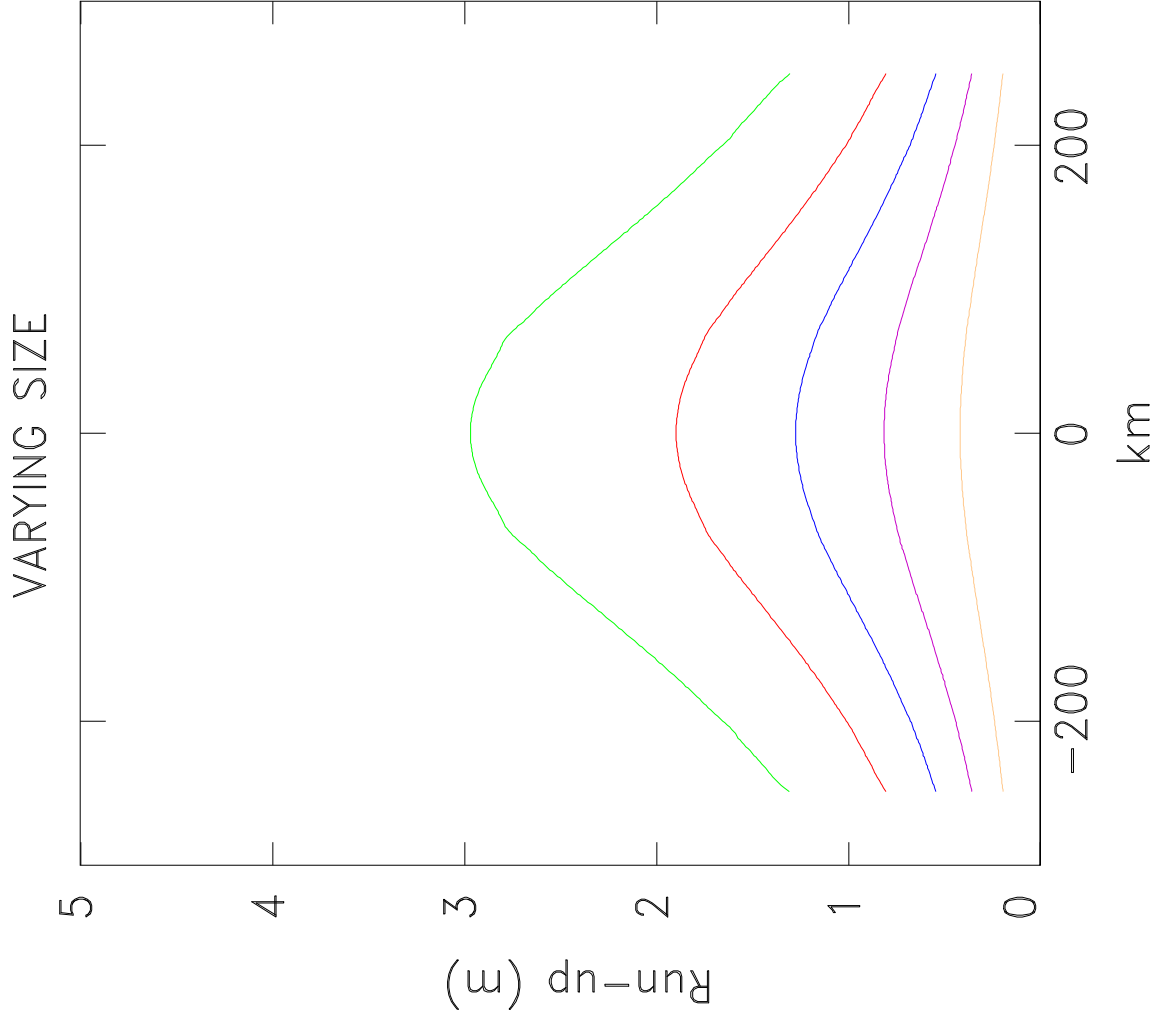
$M_0 = 2. \cdot 10^{28} \text{ dyn-cm}$

Slip on fault

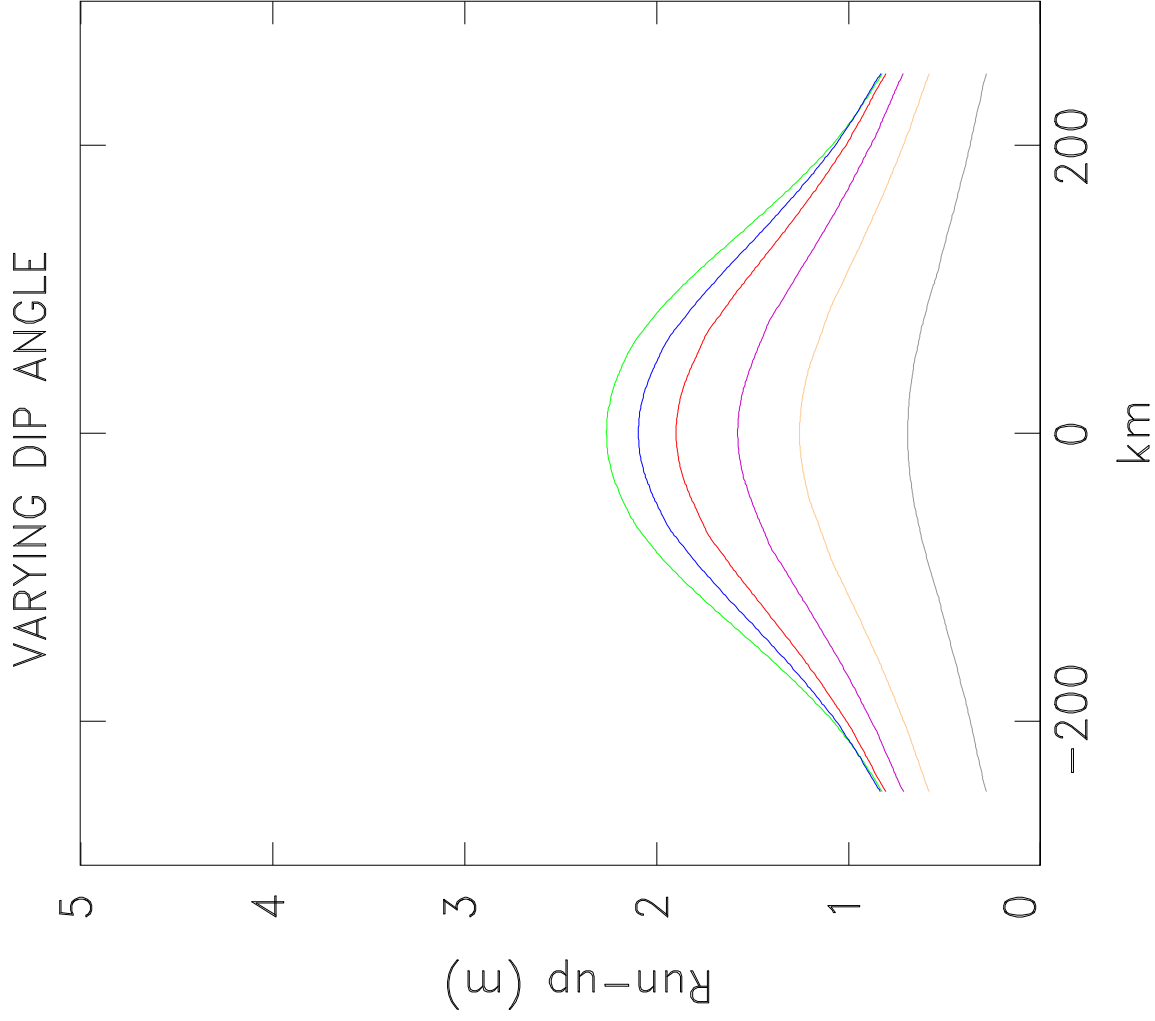
$\Delta u = 4.13 \text{ m}$

$I_1 = b/\Delta u = 0.46$

- Vary source parameters: I no greater than 2.3×10^{-5} .

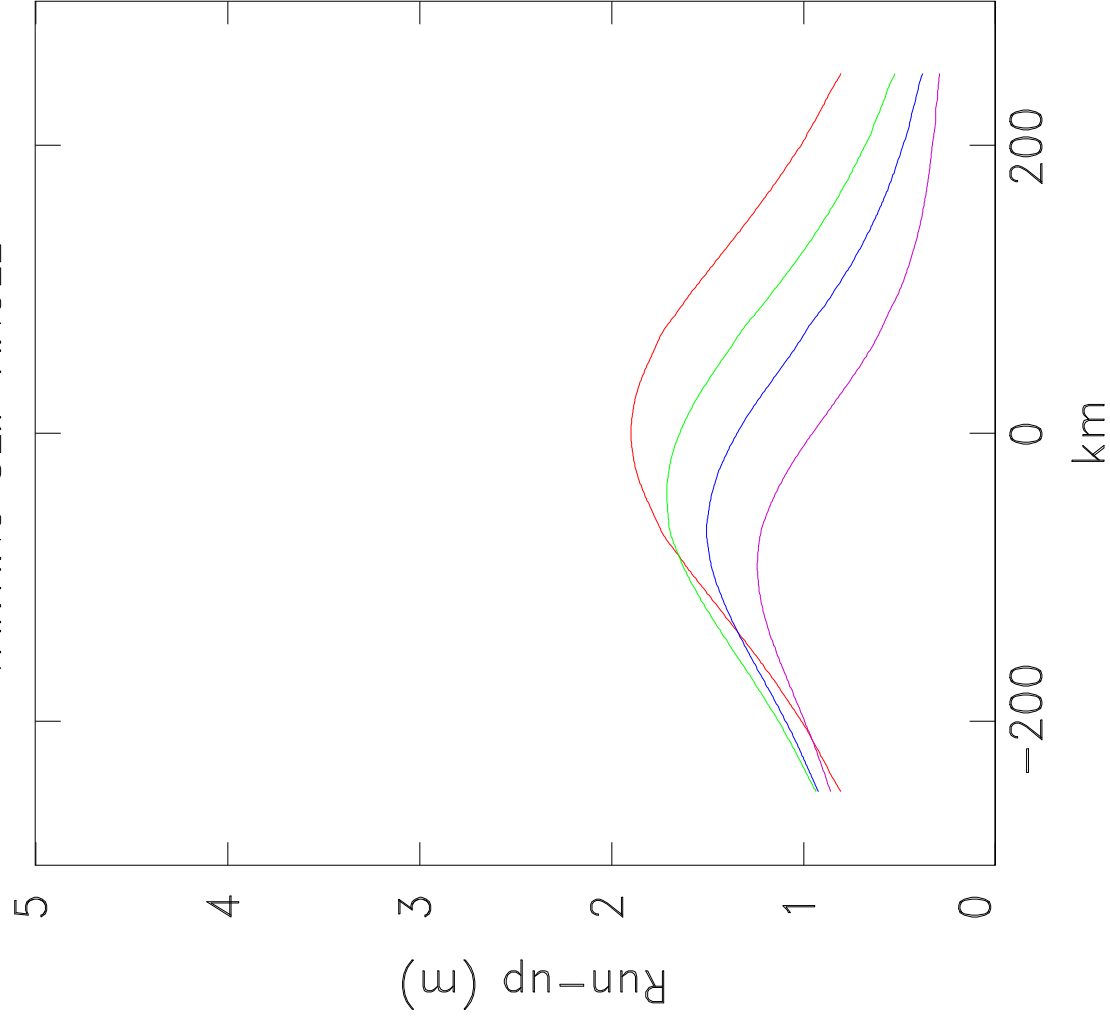


INDEX	MOMENT 10 ⁻²⁷ dyn-cm	Aspect Ratio (10 ^{**} -5)	b/ Δu
124	20.	0.86	0.45
135	50.	1.30	0.53
136	10.	0.60	0.39
137	5.	0.37	0.32
138	2.	0.18	0.22



INDEX	DIP (deg.)	Aspect Ratio (10** ⁻⁵)	b/ Δu
124	30.	0.86	0.45
126	10.	1.20	0.56
125	20.	1.04	0.51
123	45.	0.70	0.38
122	60.	0.54	0.30
121	90.	0.33	0.17

VARYING SLIP ANGLE



INDEX

SLIP (deg.)

Aspect Ratio (10**⁻⁵)

b/ Δu

124

90.

0.86

0.45

127

60.

0.83

0.42

128

45.

0.77

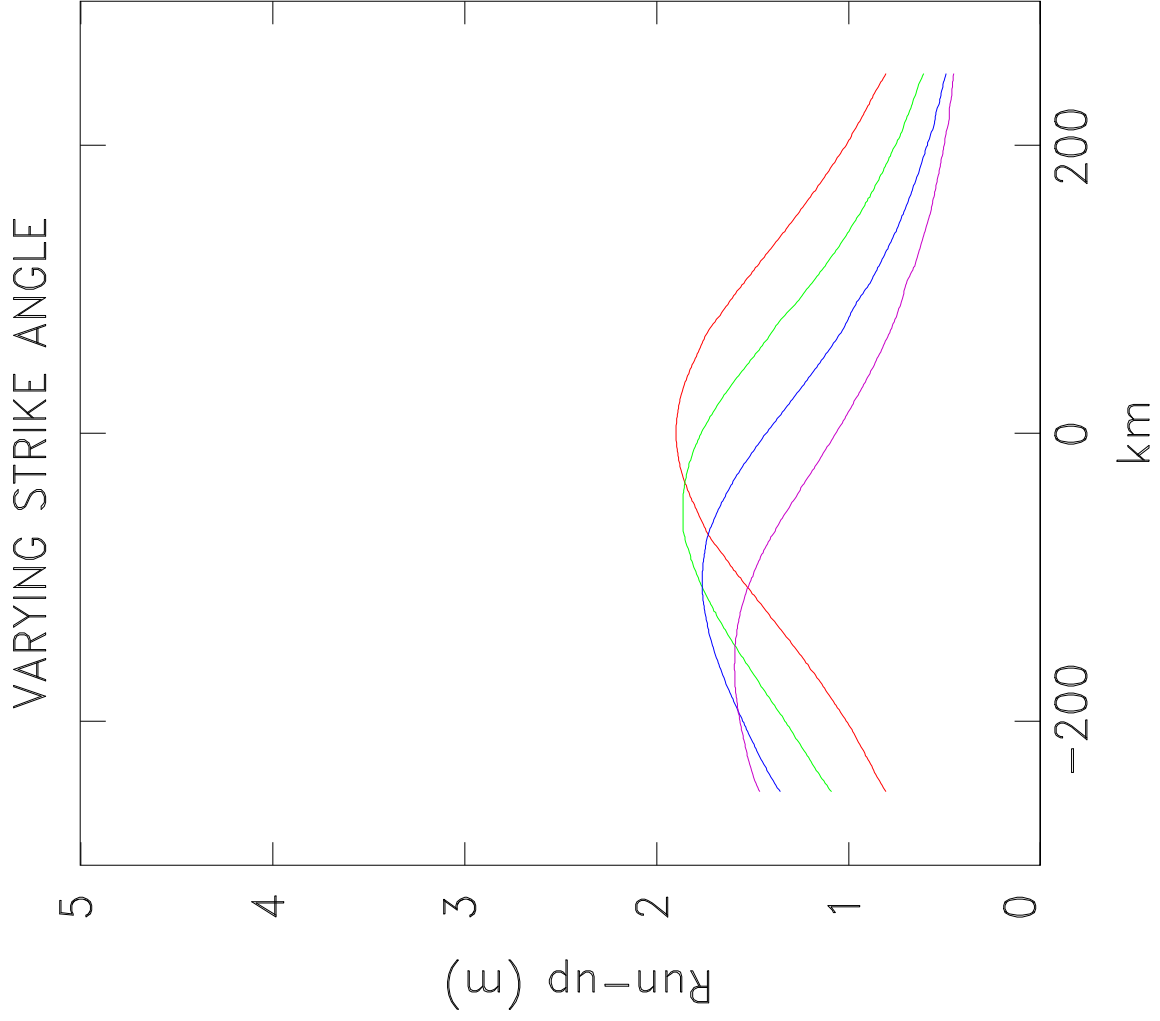
0.37

129

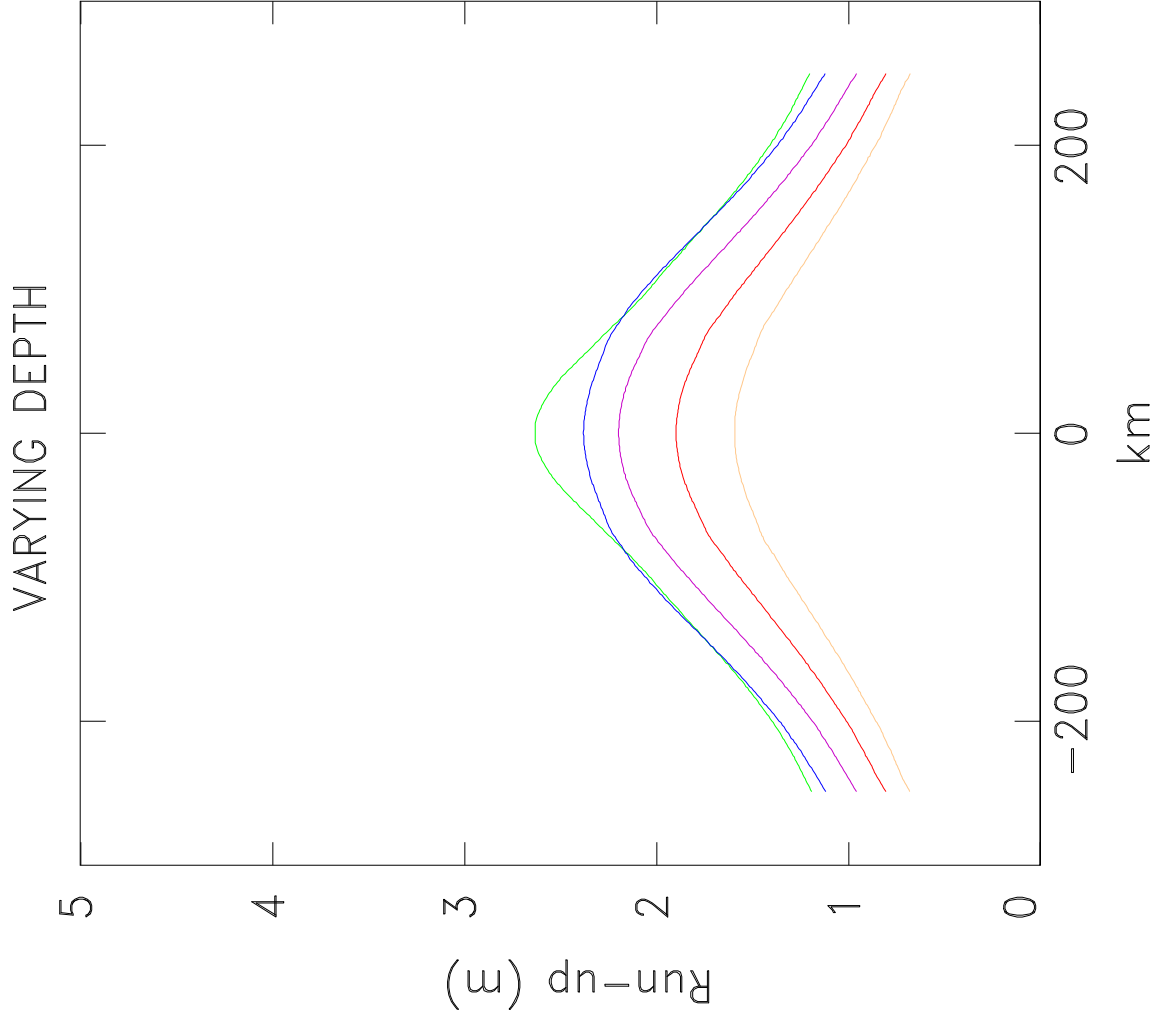
30.

0.70

0.31

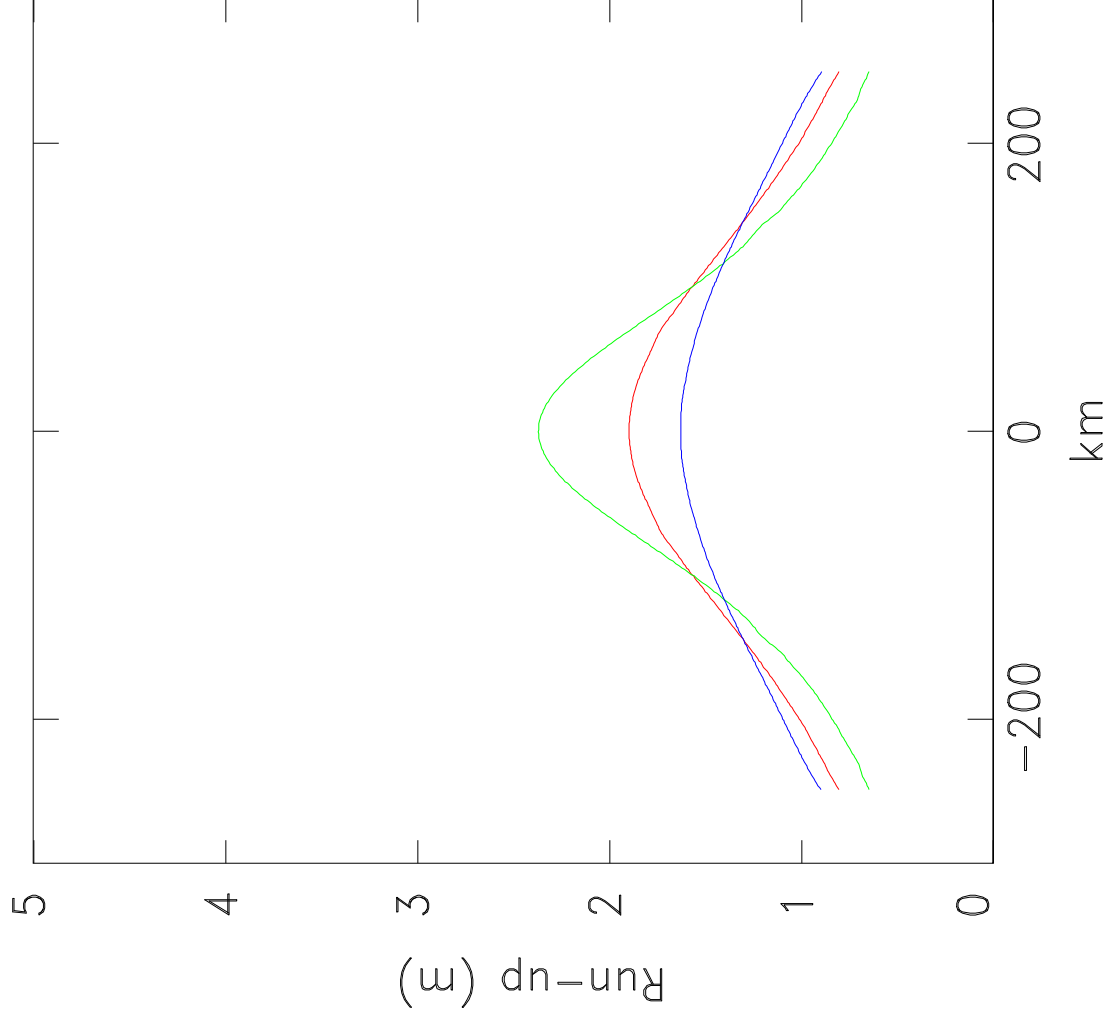


INDEX	STRIKE (deg.)	Aspect Ratio (10^{*-5})	$b/\Delta u$
124	270.	0.86	0.45
141	285.	0.86	0.45
139	300.	0.80	0.43
140	315.	0.67	0.28



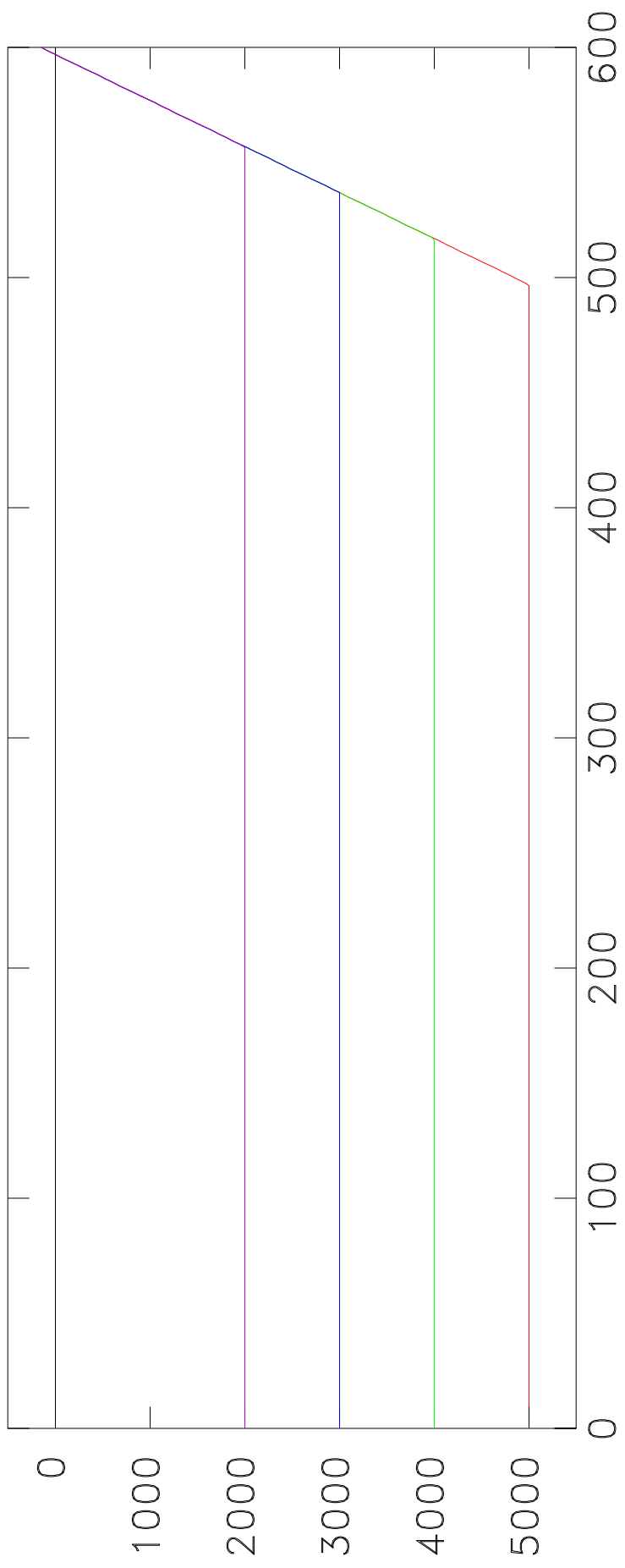
INDEX	DEPTH (km)	Aspect Ratio (10^{*-5})	$b/\Delta u$
124	35.	0.86	0.45
131	5.	1.20	0.62
132	15.	1.04	0.58
133	25.	1.01	0.54
134	45.	0.74	0.38

VARYING DISTANCE from BEACH

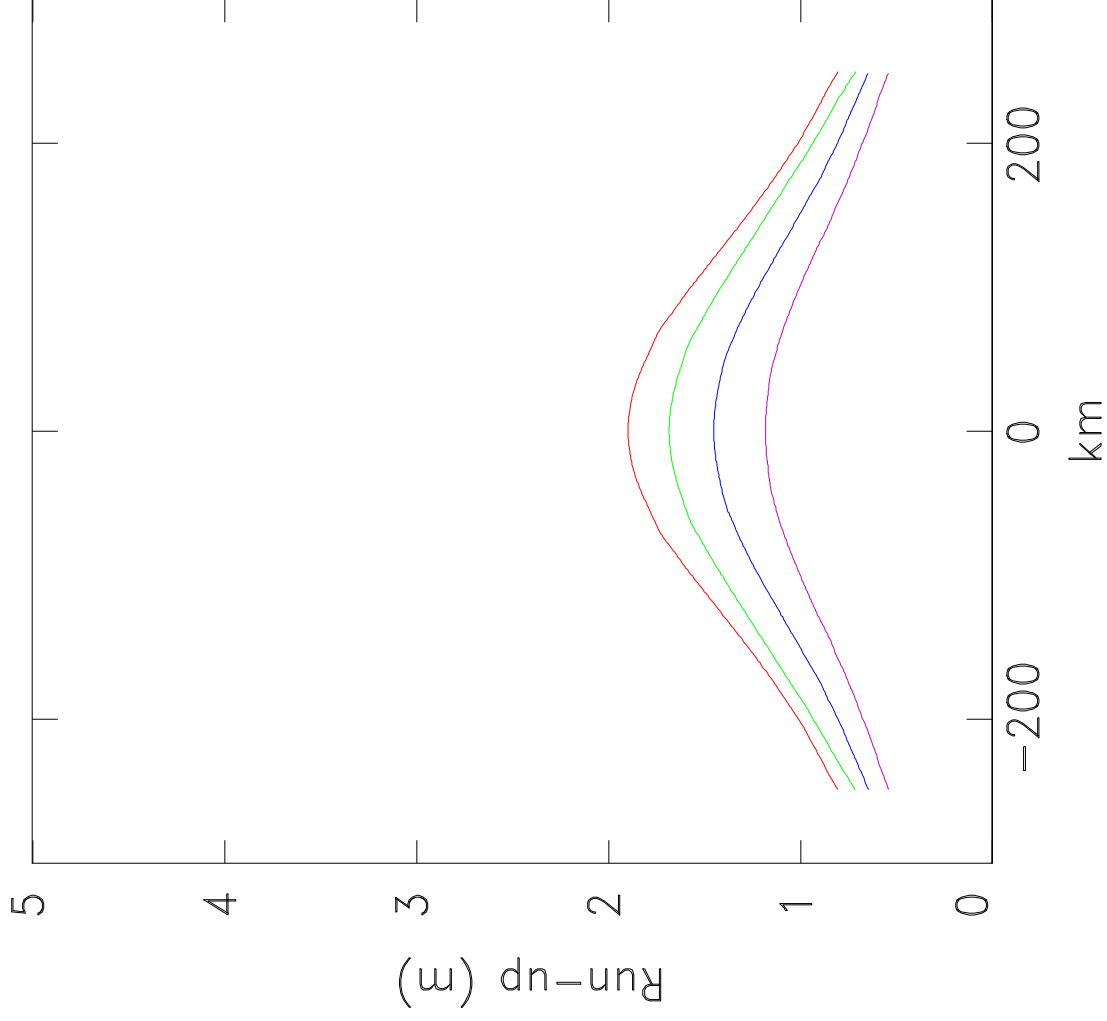


INDEX	INCREM. DIST (km)	Aspect Ratio (10^{*-5})	$b/\Delta u$
124	0.	0.86	0.45
143	-100.	1.59	0.57
144	100.	2.90	0.69

MODELS 124, 145, 146, 147



VARYING WATER DEPTH



INDEX

WATER DEPTH

(m)

Aspect Ratio

(10**⁻⁵)

b/ Δ u

124

5000.

0.86

0.45

145

4000.

0.76

0.41

146

3000.

0.65

0.35

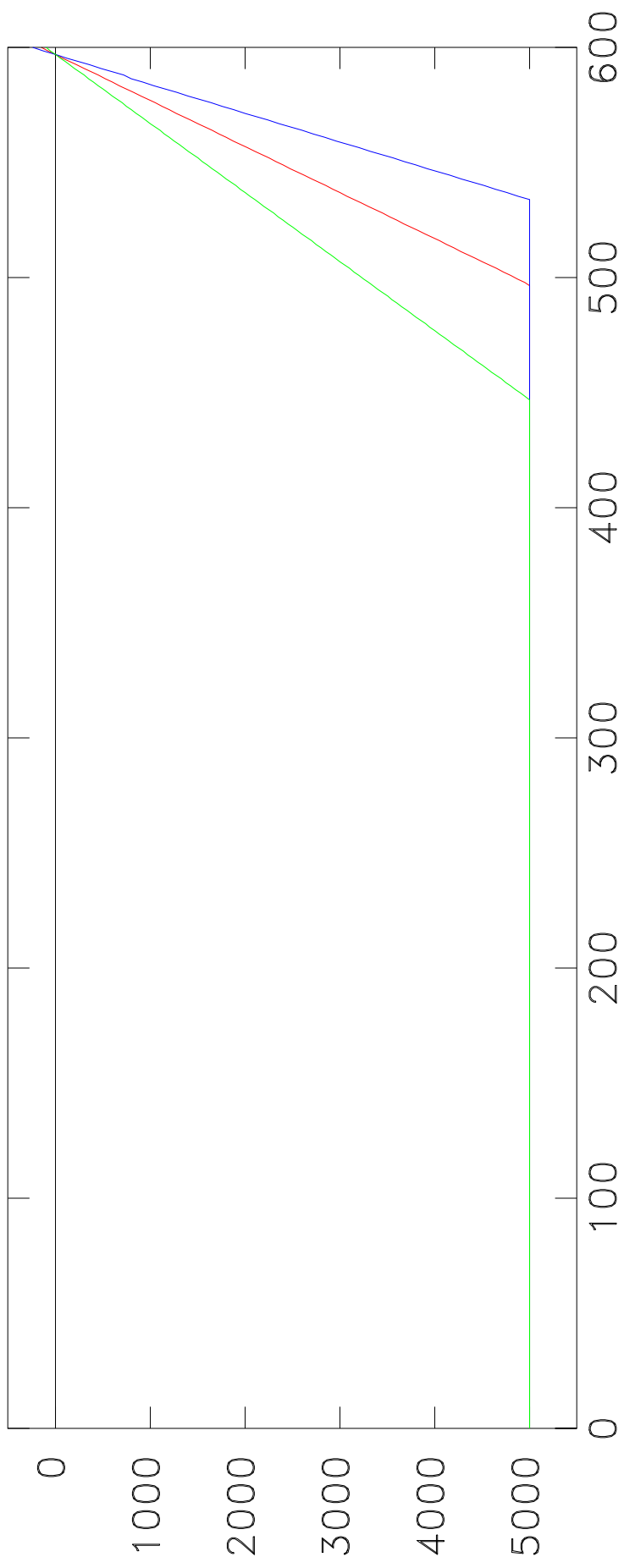
147

2000.

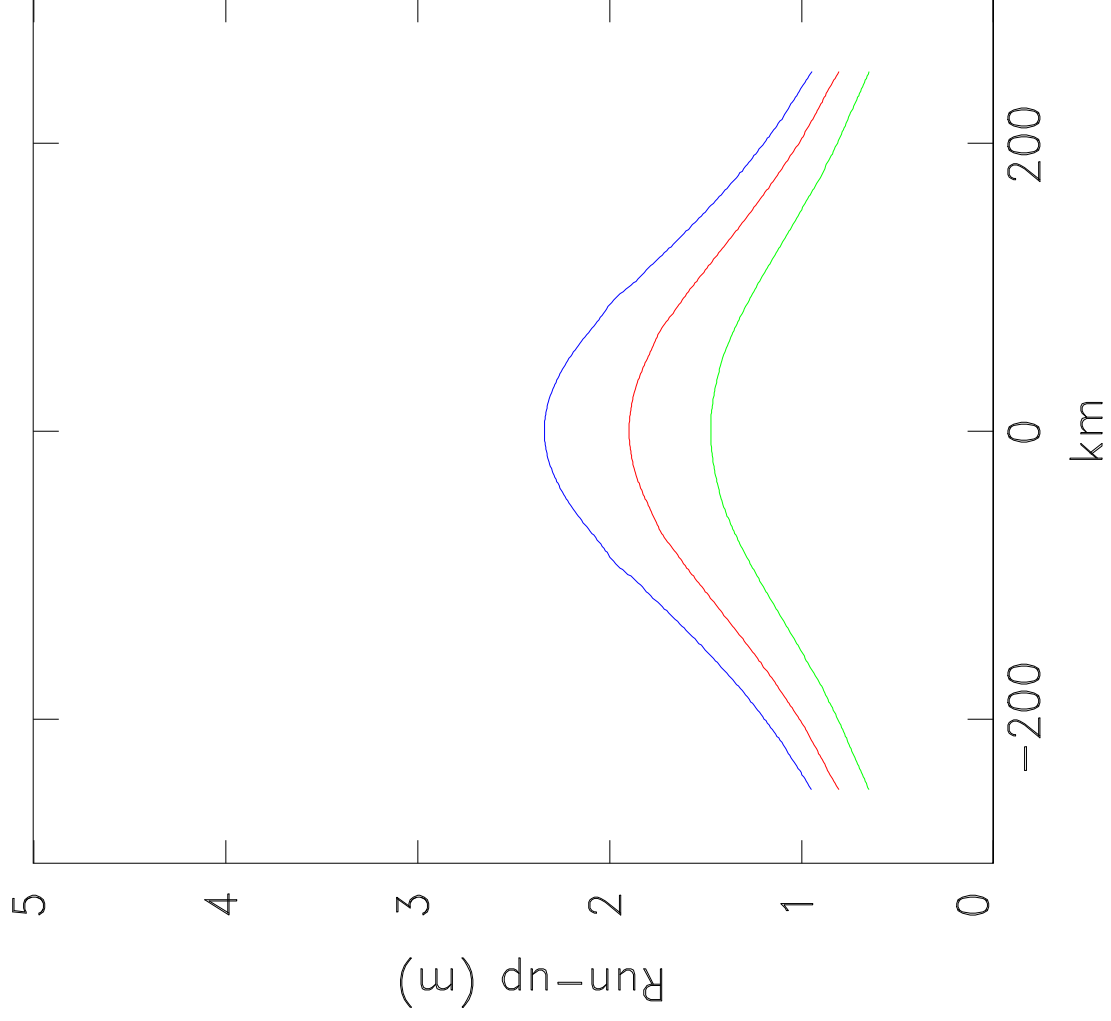
0.52

0.29

MODELS 124, 148, 149



VARYIGN BEACH SLOPE



INDEX

SLOPE
(m/km)

Aspect Ratio
(10**⁻⁵)

b/Δu

124

50.

0.86

0.45

148

80.

0.68

0.36

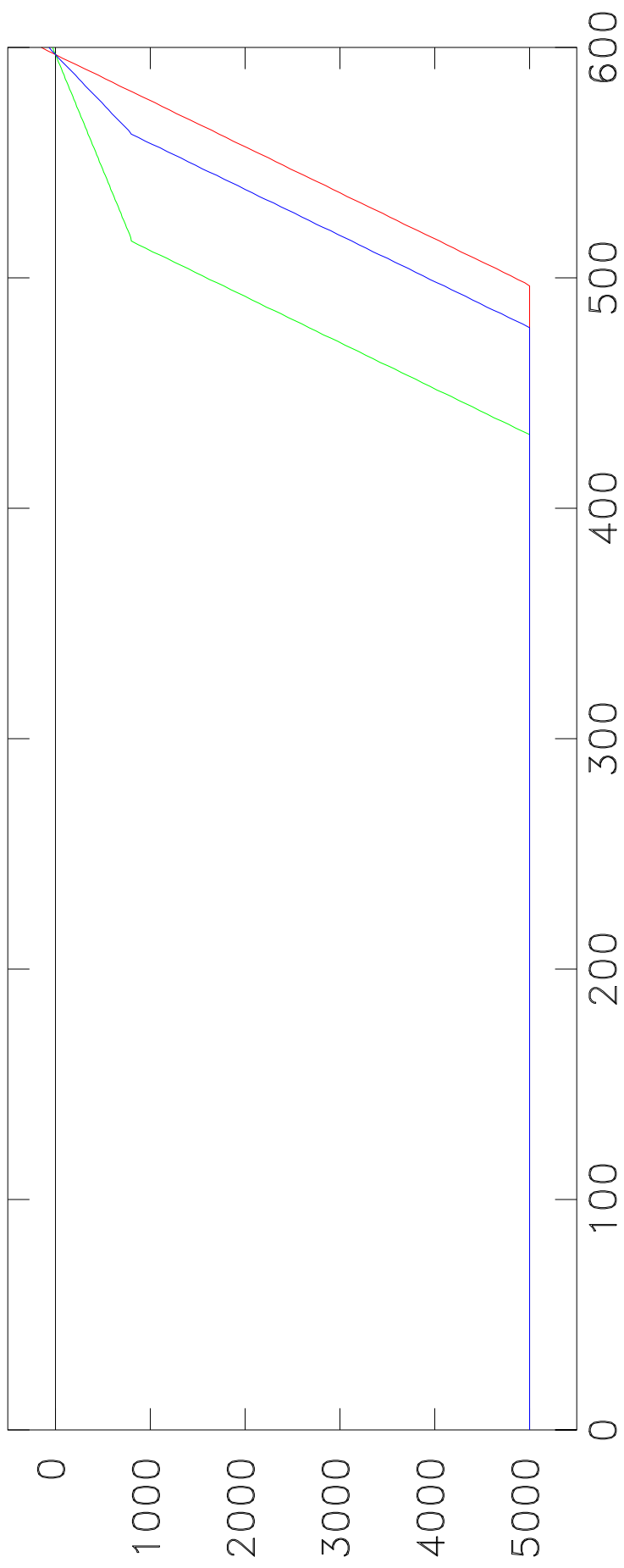
149

33.

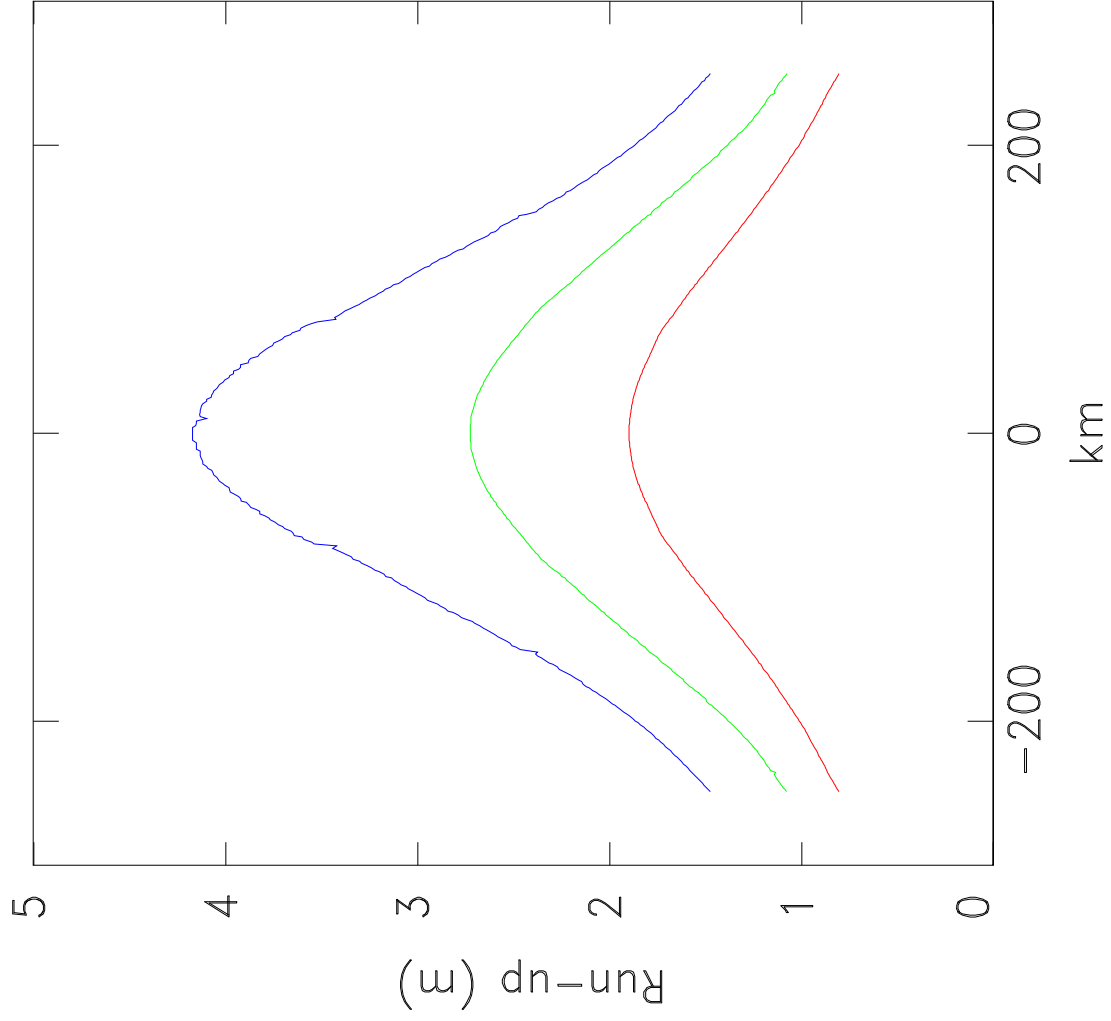
1.12

0.55

MODELS 124, 150, 151



COMPOSITE BEACHES



INDEX

BEACH No.

b/ Δu

Aspect Ratio
(10**(-5))

124

0.

0.86

0.45

151

1.

1.30

0.65

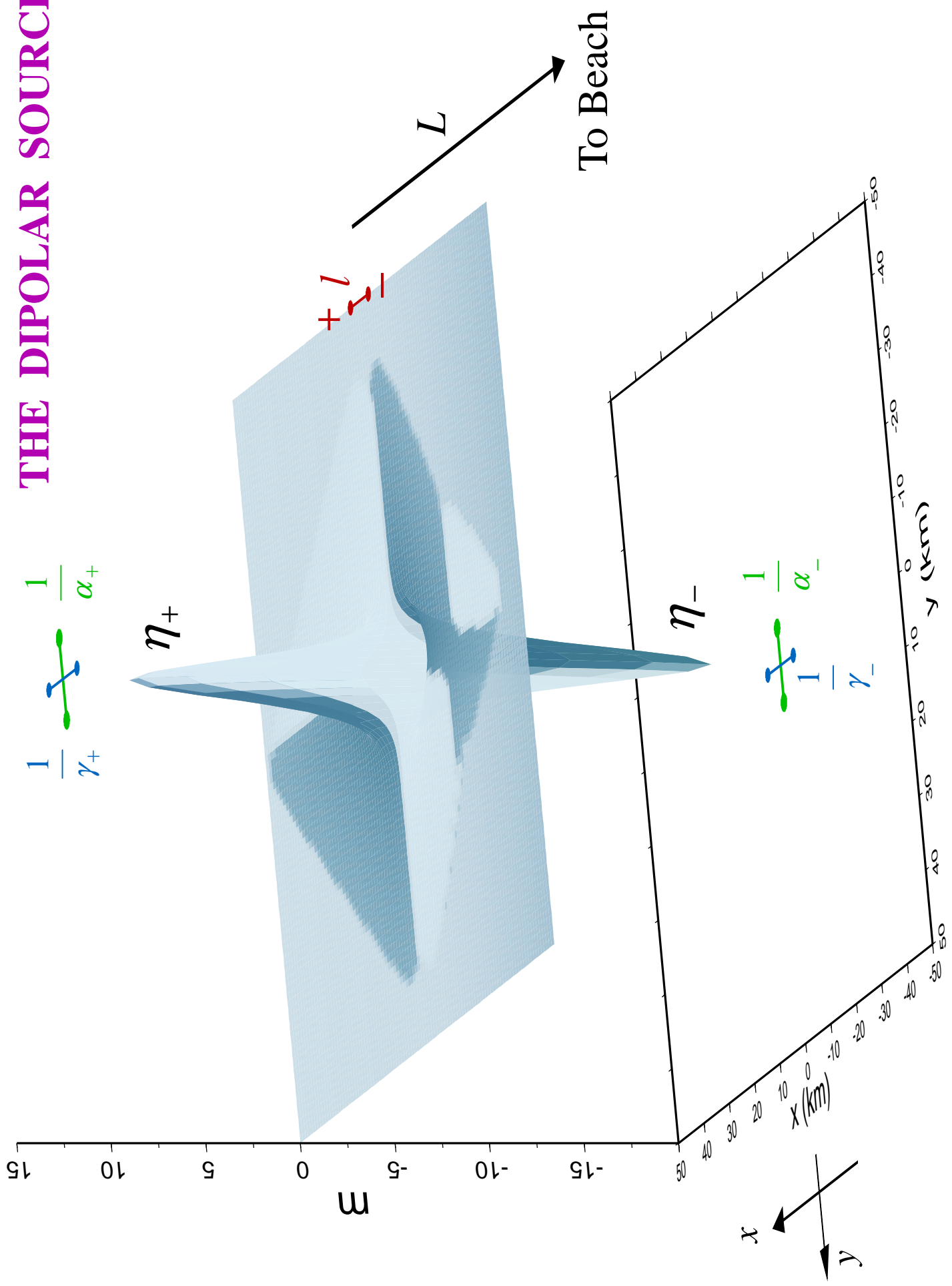
150

2.

2.30

0.80

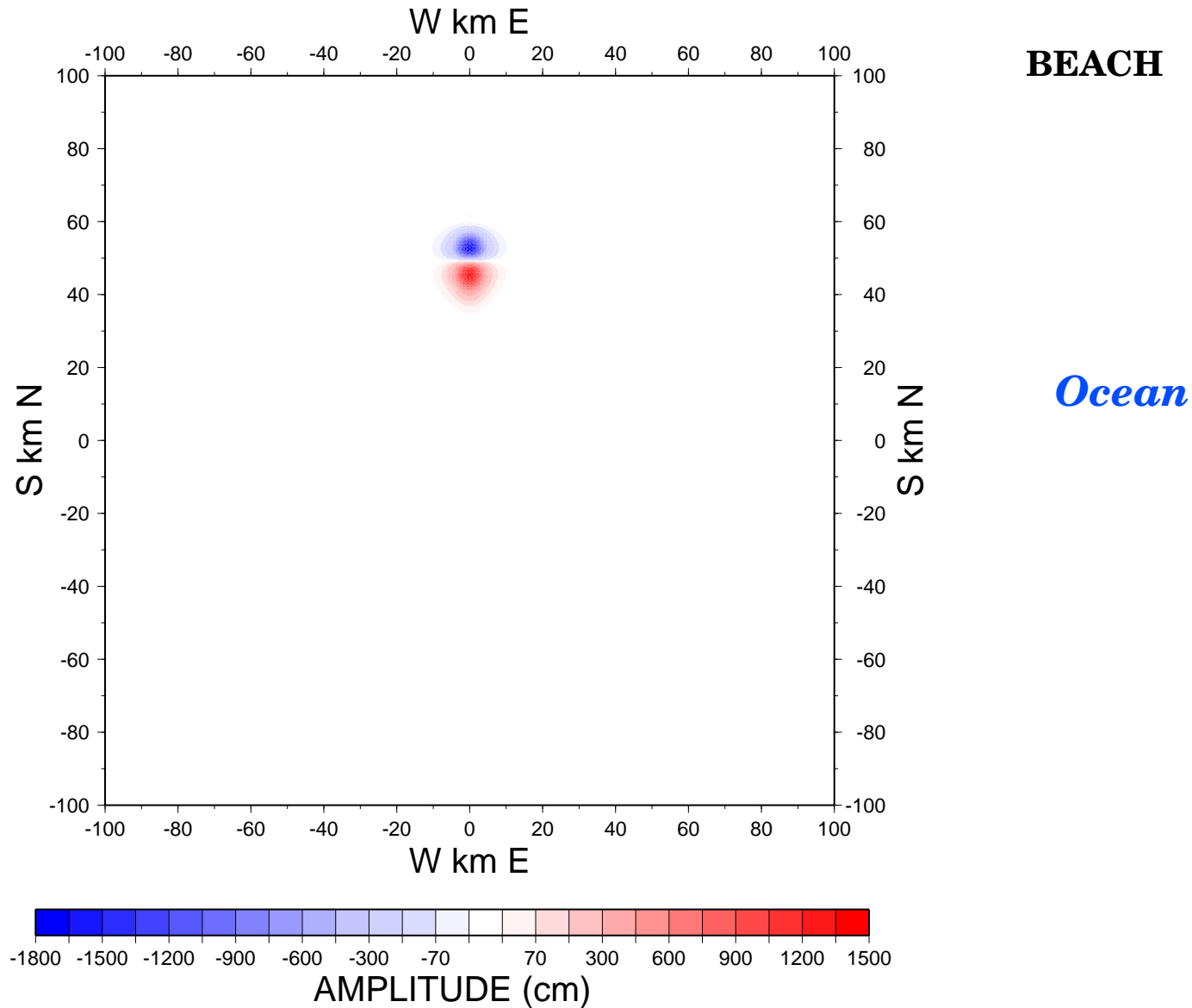
THE DIPOLAR SOURCE



[Okal and Synolakis, in press]

NEAR-FIELD: *The Landslide Source*

- Compute Ocean-Surface Deformation due to Landslide

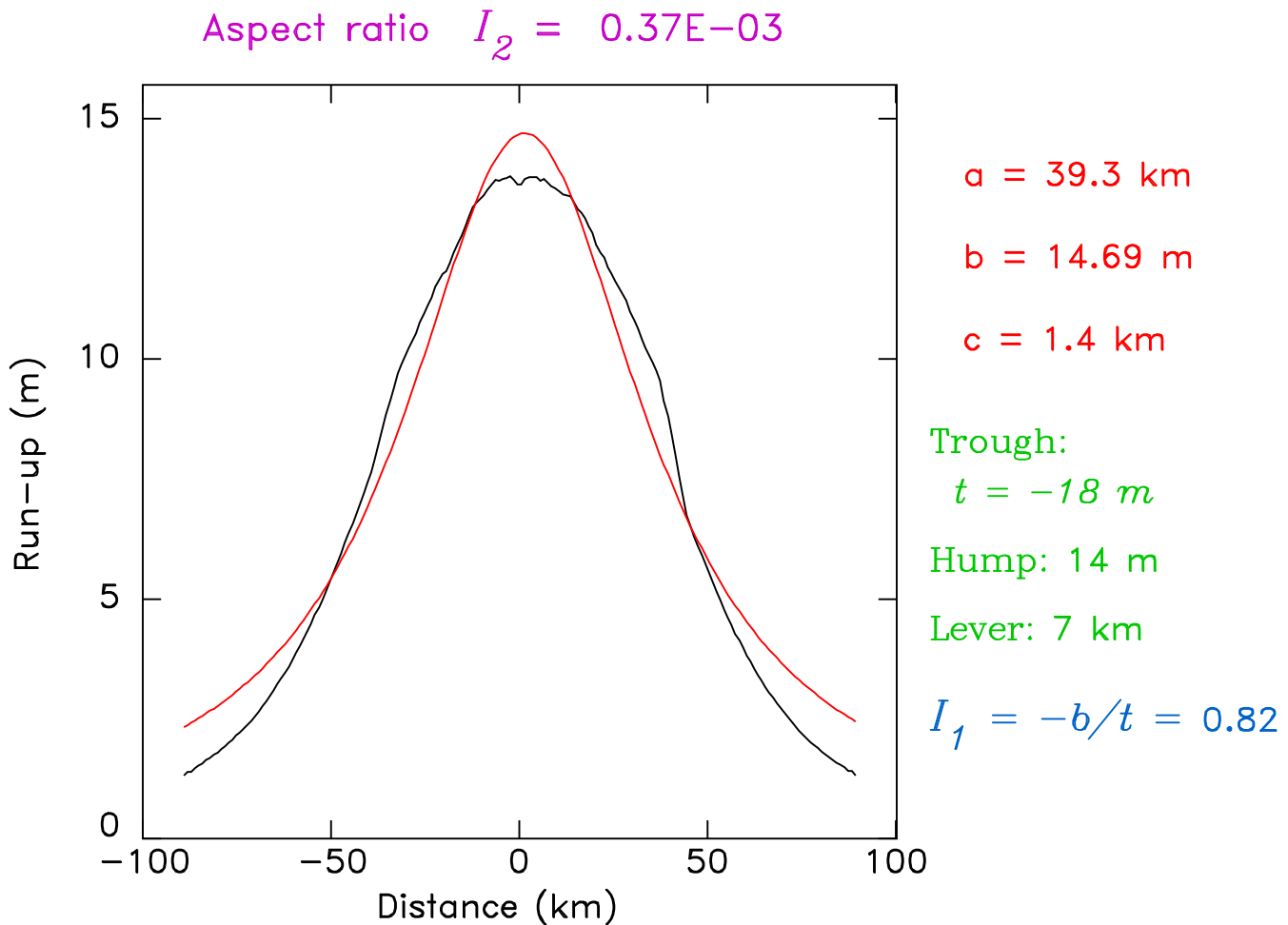


- Simulate Tsunami Propagation to Beach and Run-up

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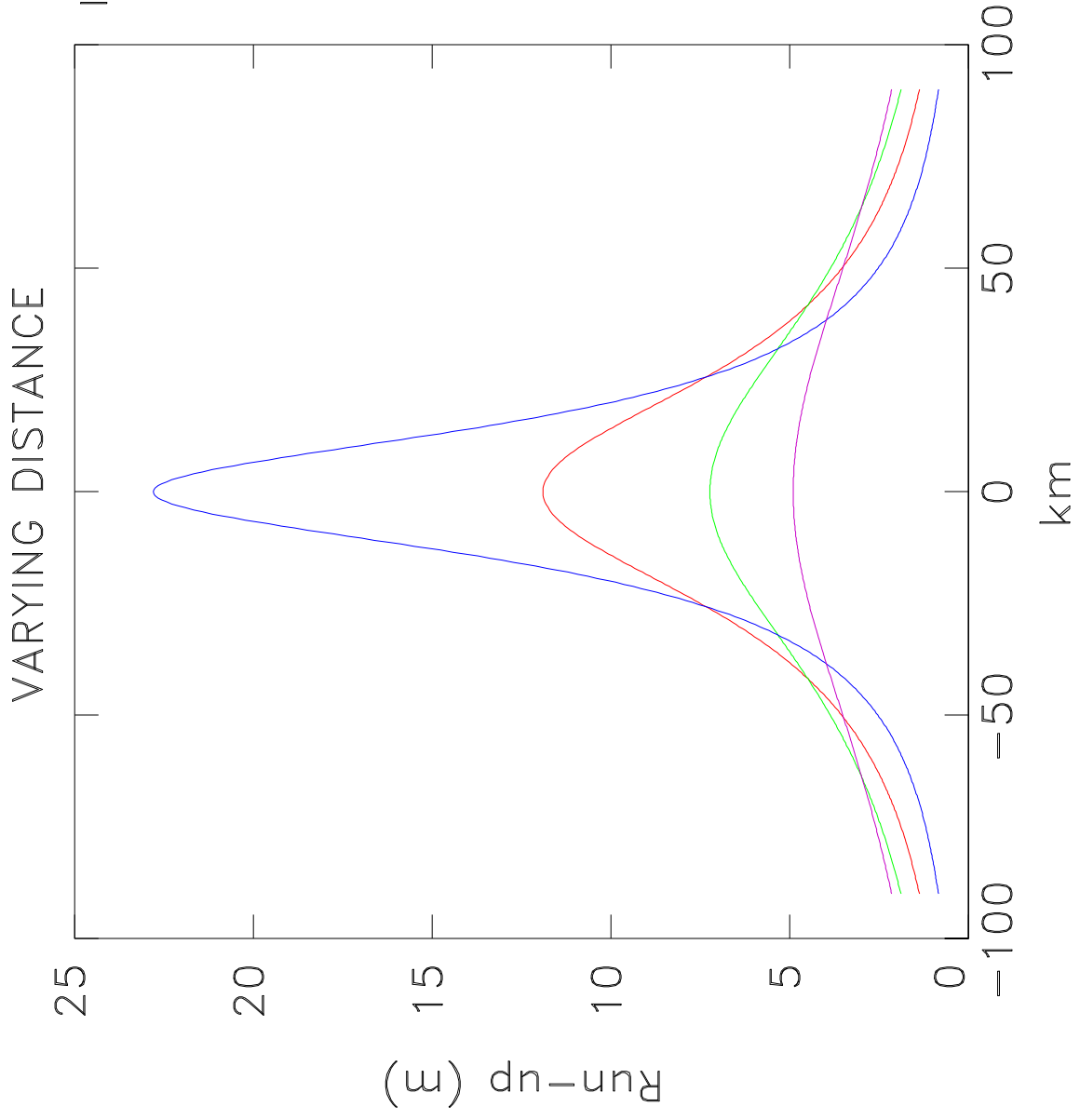
- Fit Bell Curve
$$\zeta = \frac{b}{\left(\frac{x - c}{a}\right)^2 + 1}$$

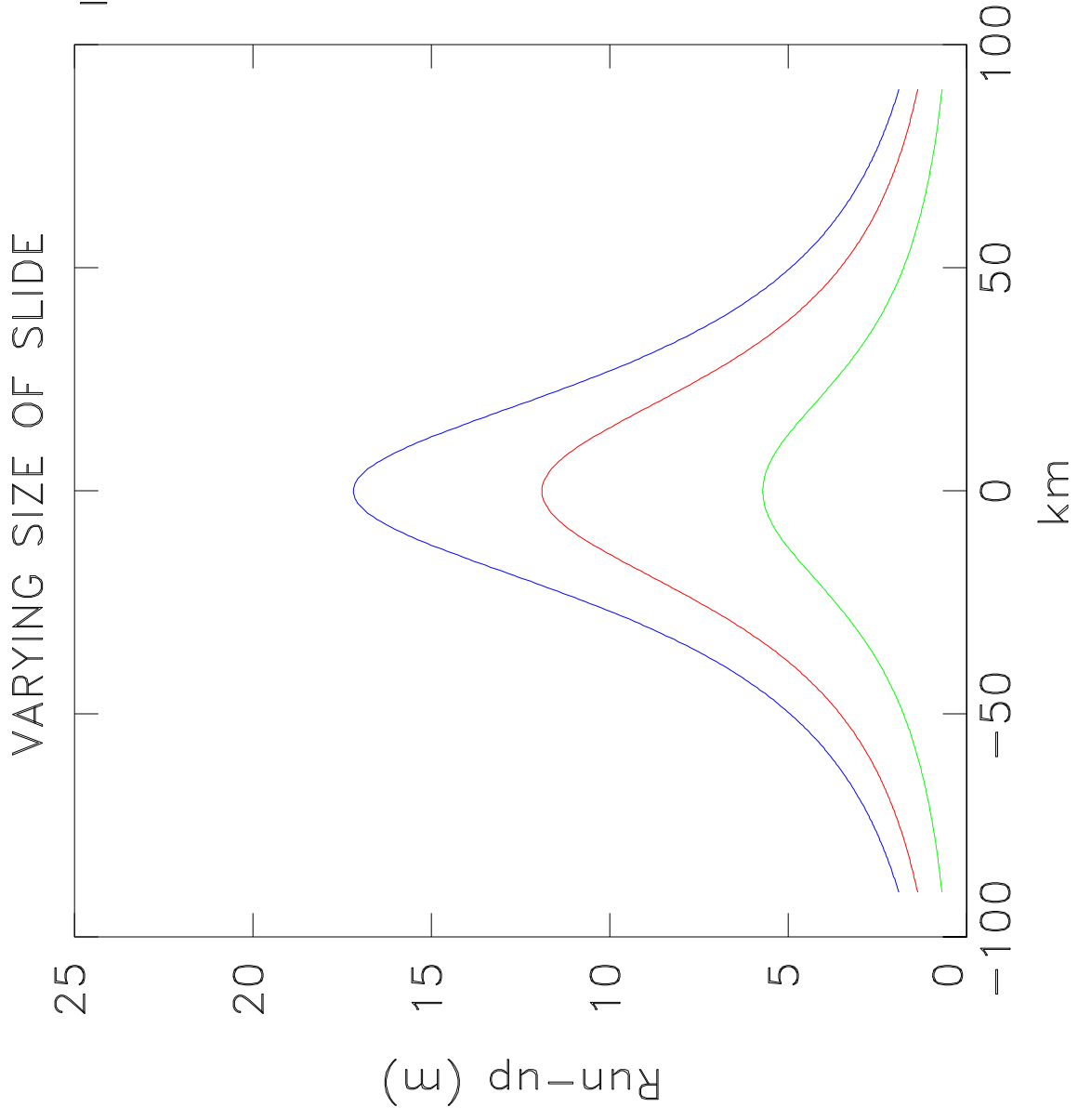
- Retain aspect ratio $I = b/a$



- Vary source parameters: I greater than 10^{-4} .

$I = b/a$ CAN SERVE AS DISCRIMINANT



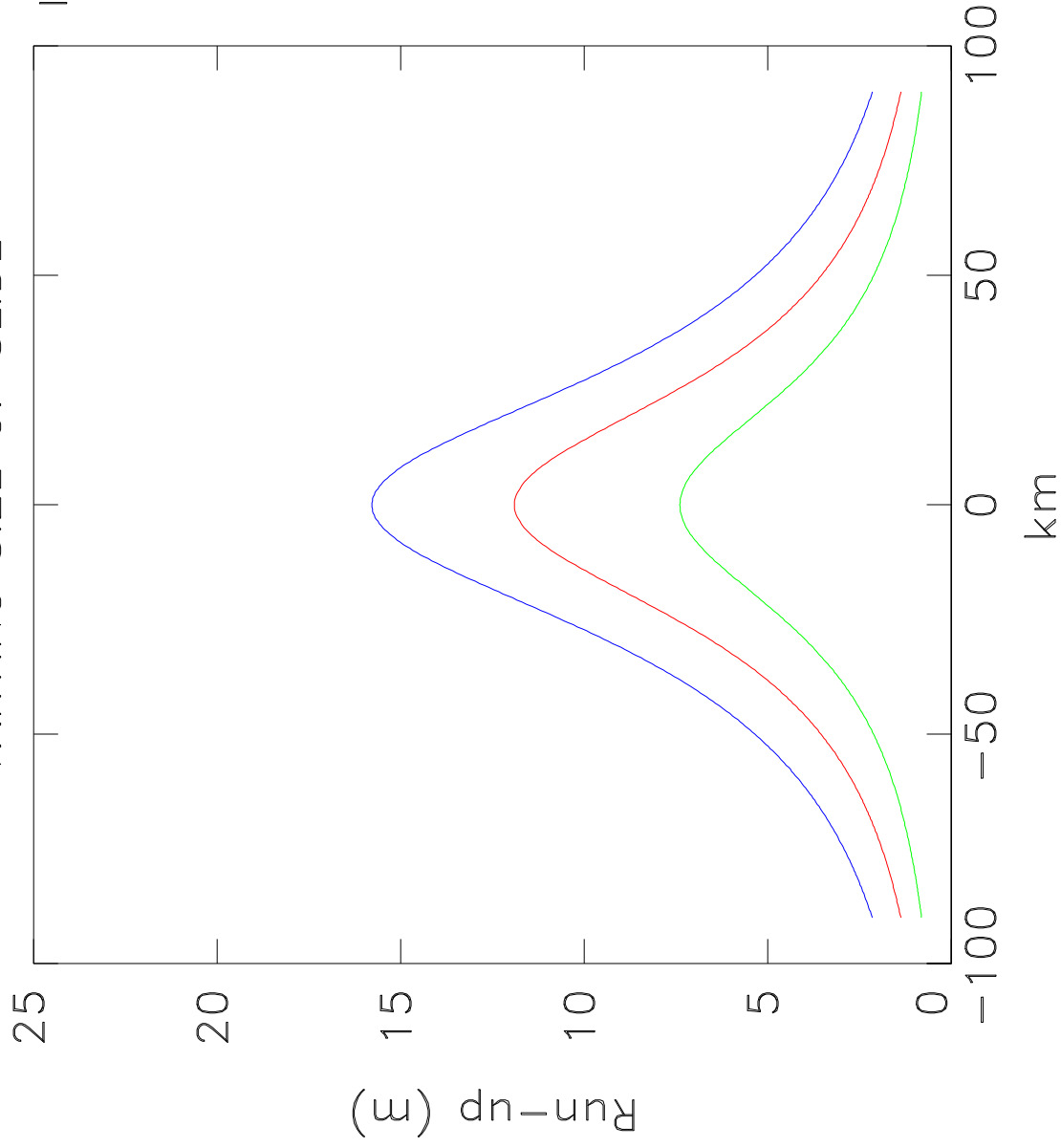


VARYING SIZE OF SLIDE

Run-up (m)

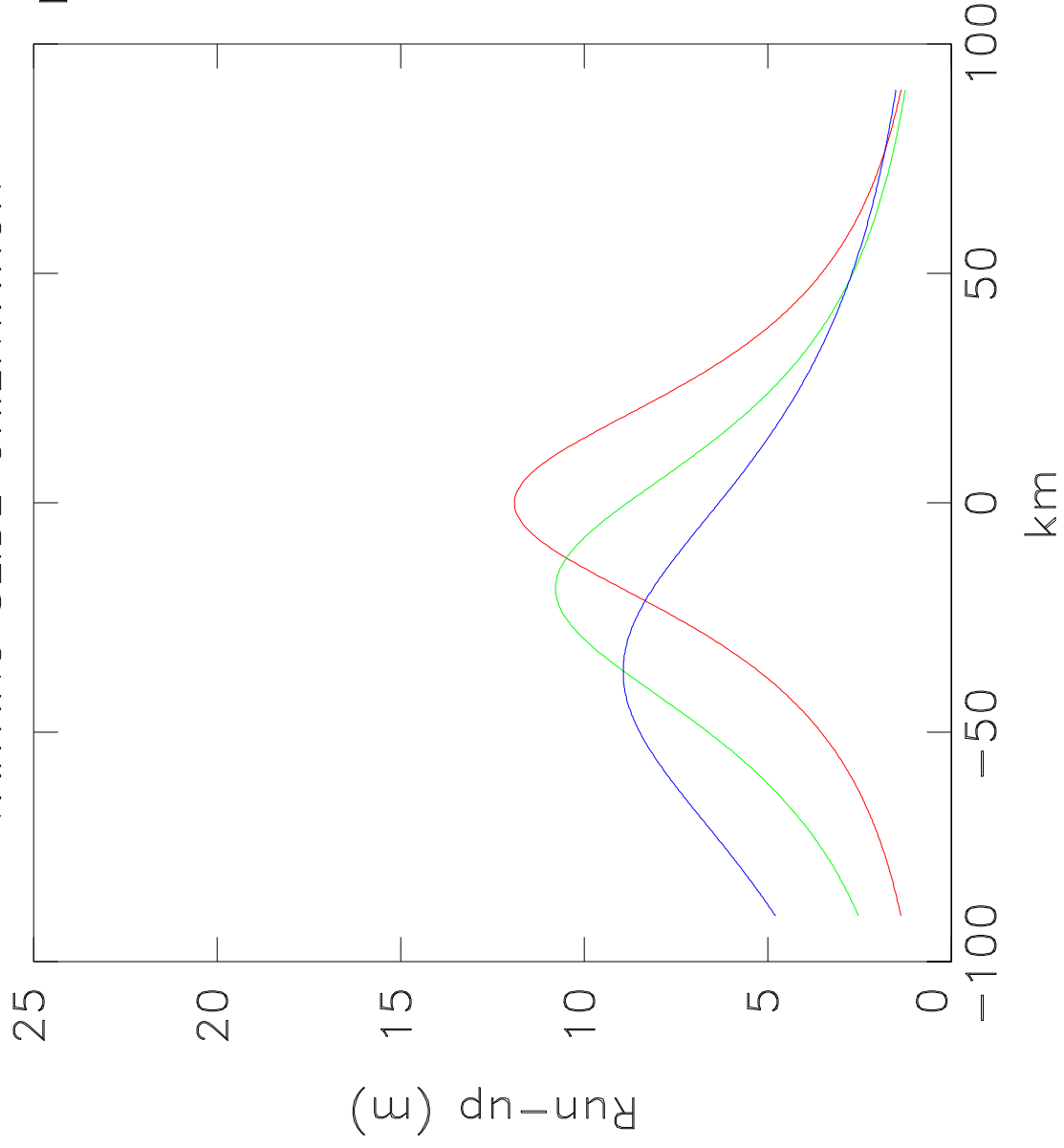
km

VARYING SIZE OF SLIDE



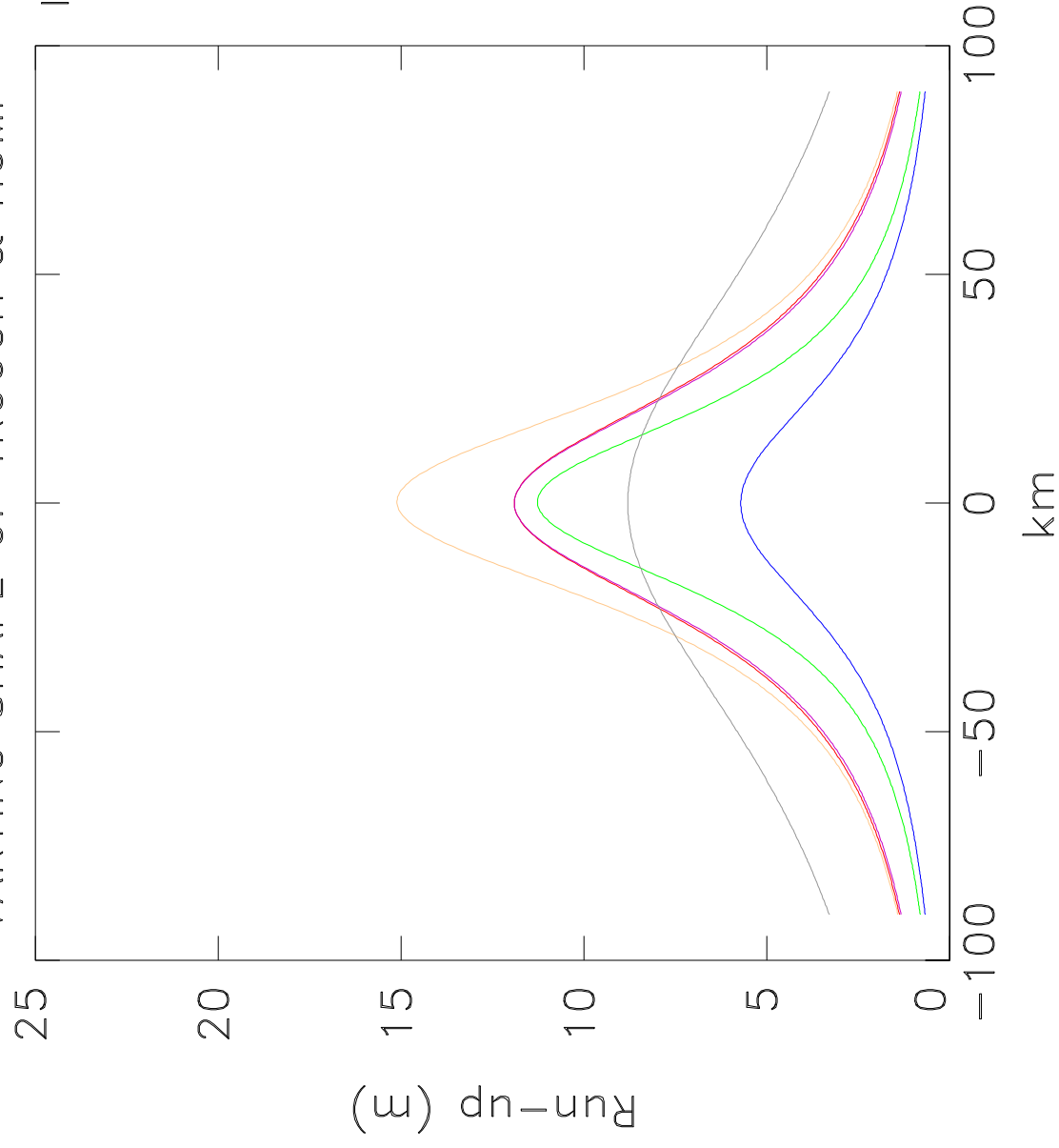
INDEX	LEVER (km)	Aspect Ratio (10**-4)
301	7.	3.65
311	4.	2.33
312	12.	4.41

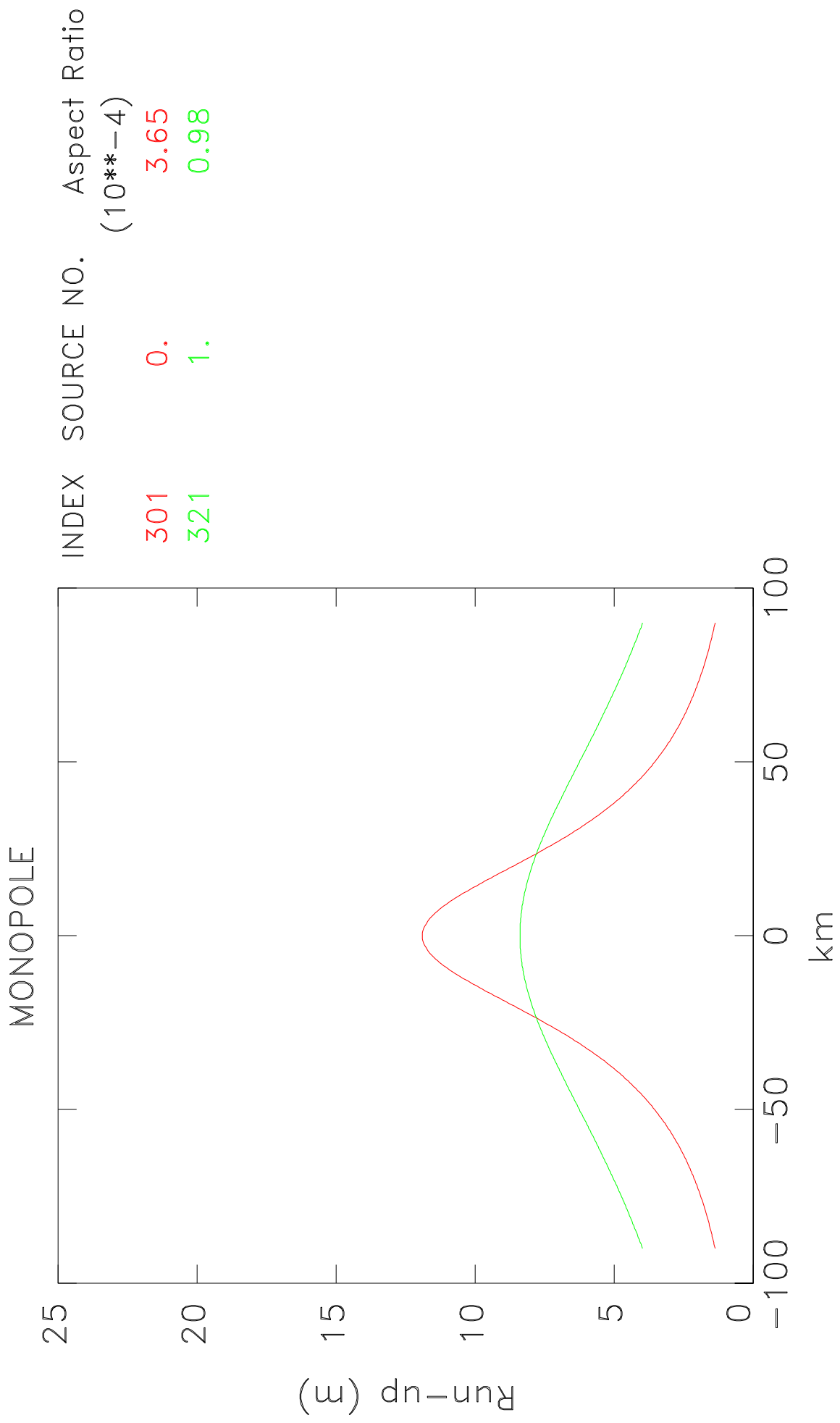
VARYING SLIDE ORIENTATION



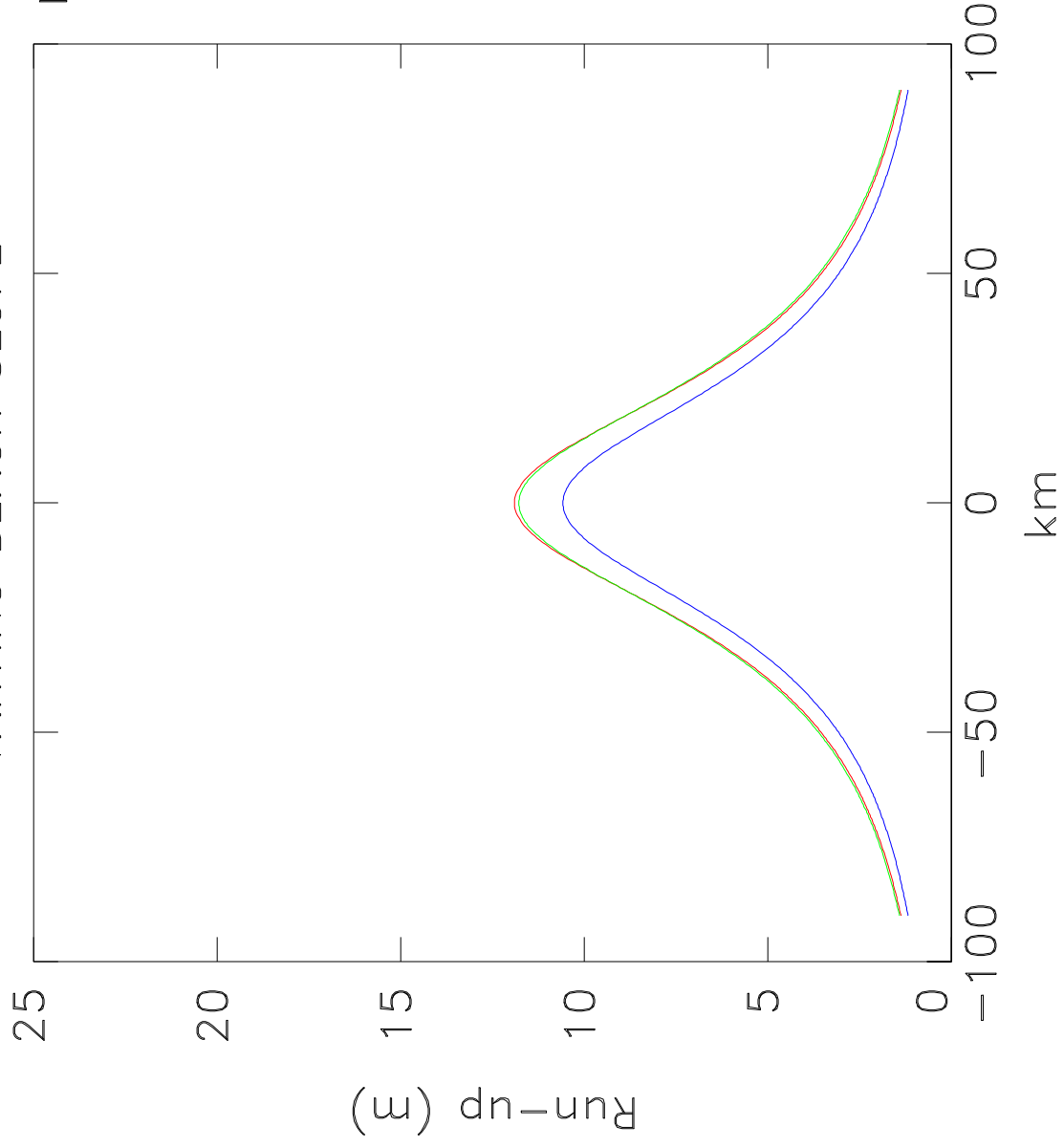
INDEX	AZIMUTH (deg.)	Aspect Ratio (10** -4)
301	180.	3.65
313	165.	2.72
312	150.	1.56

VARYING SHAPE OF TROUGH & HUMP



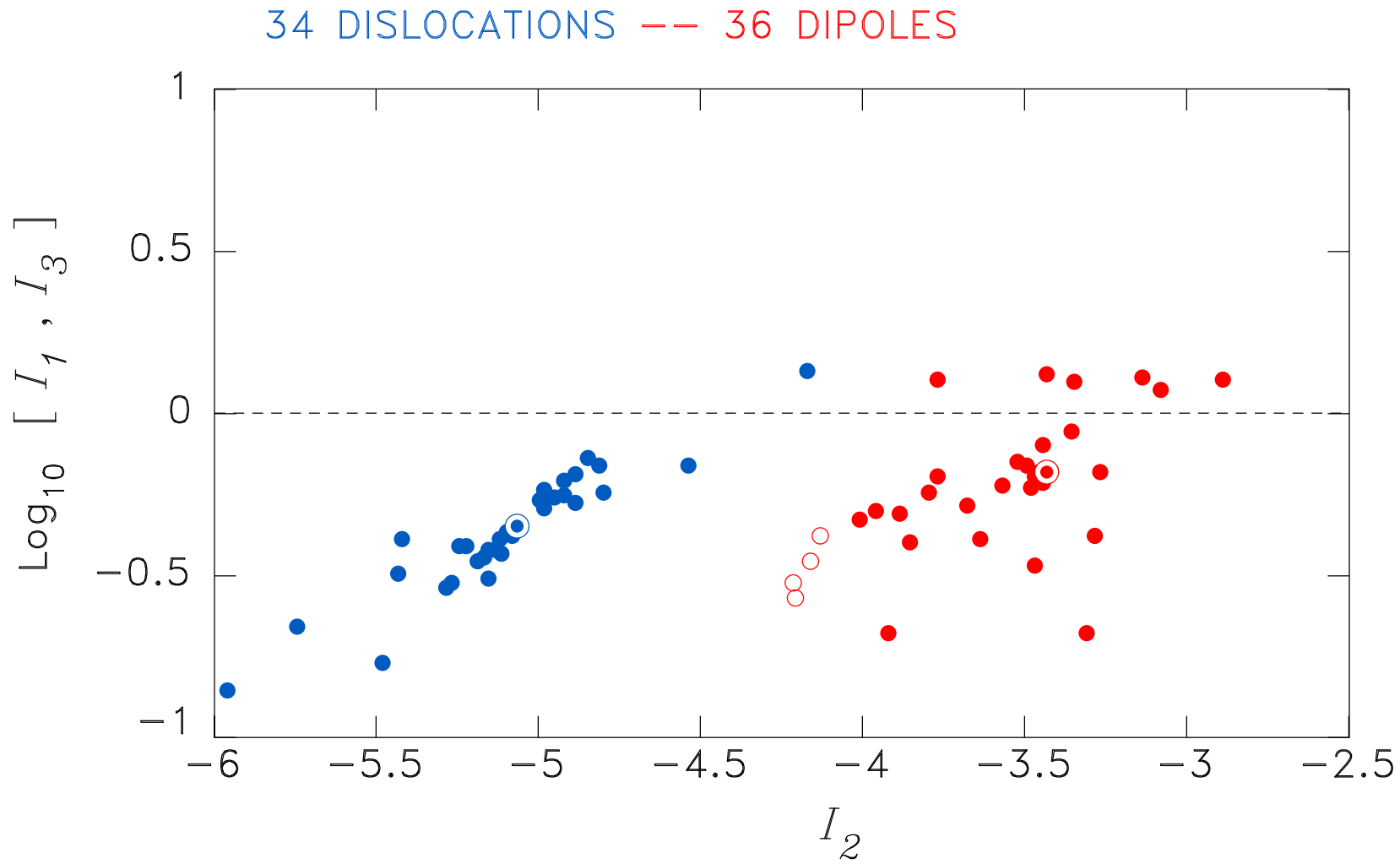


VARYING BEACH SLOPE



INDEX	SLOPE	Aspect Ratio (10**(-4))
301	0.	3.65
332	0.	3.54
330	0.	3.31

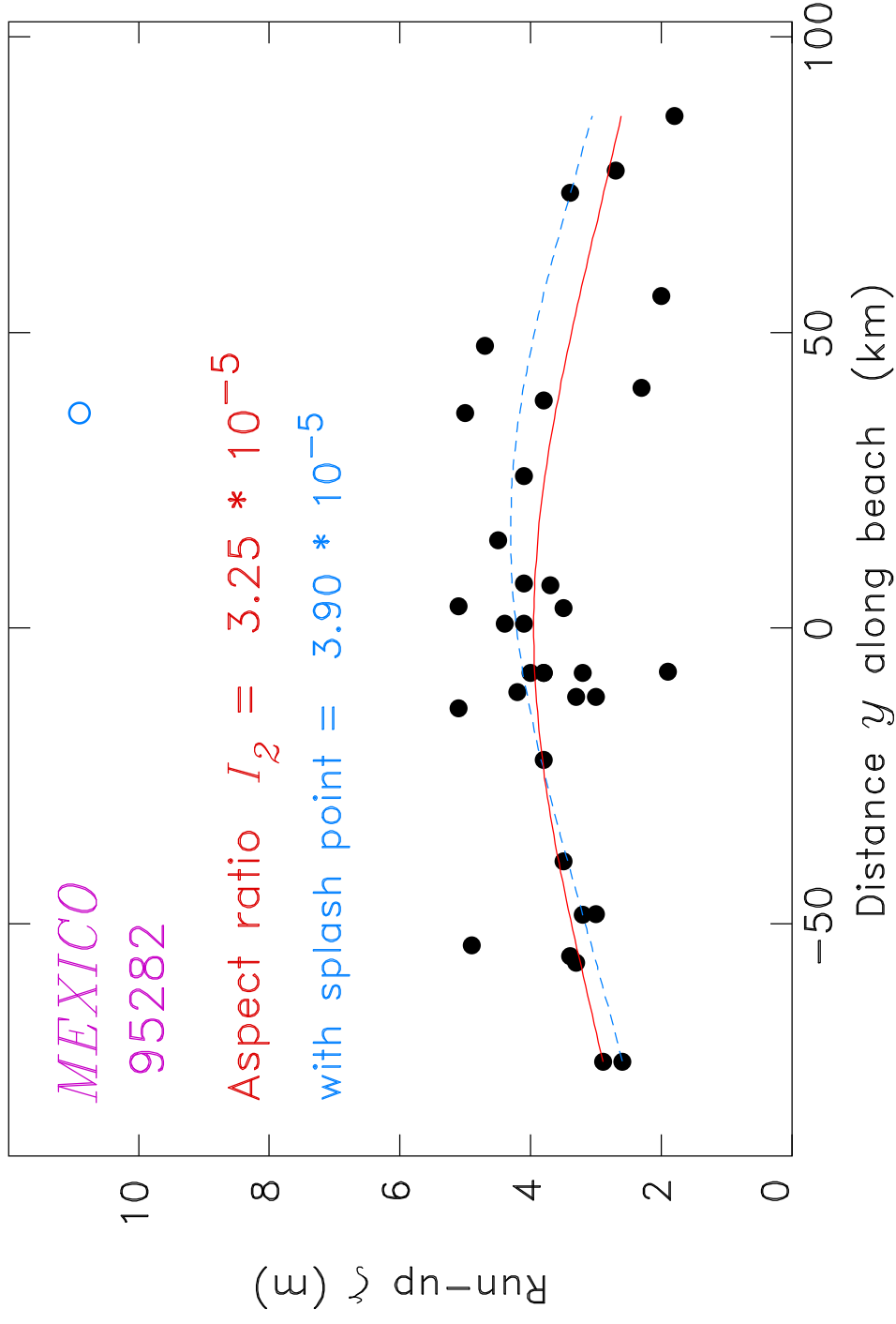
MAX. RUN-UP SCALED TO FAULT SLIP
MAX. RUN-UP SCALED TO INITIAL TROUGH



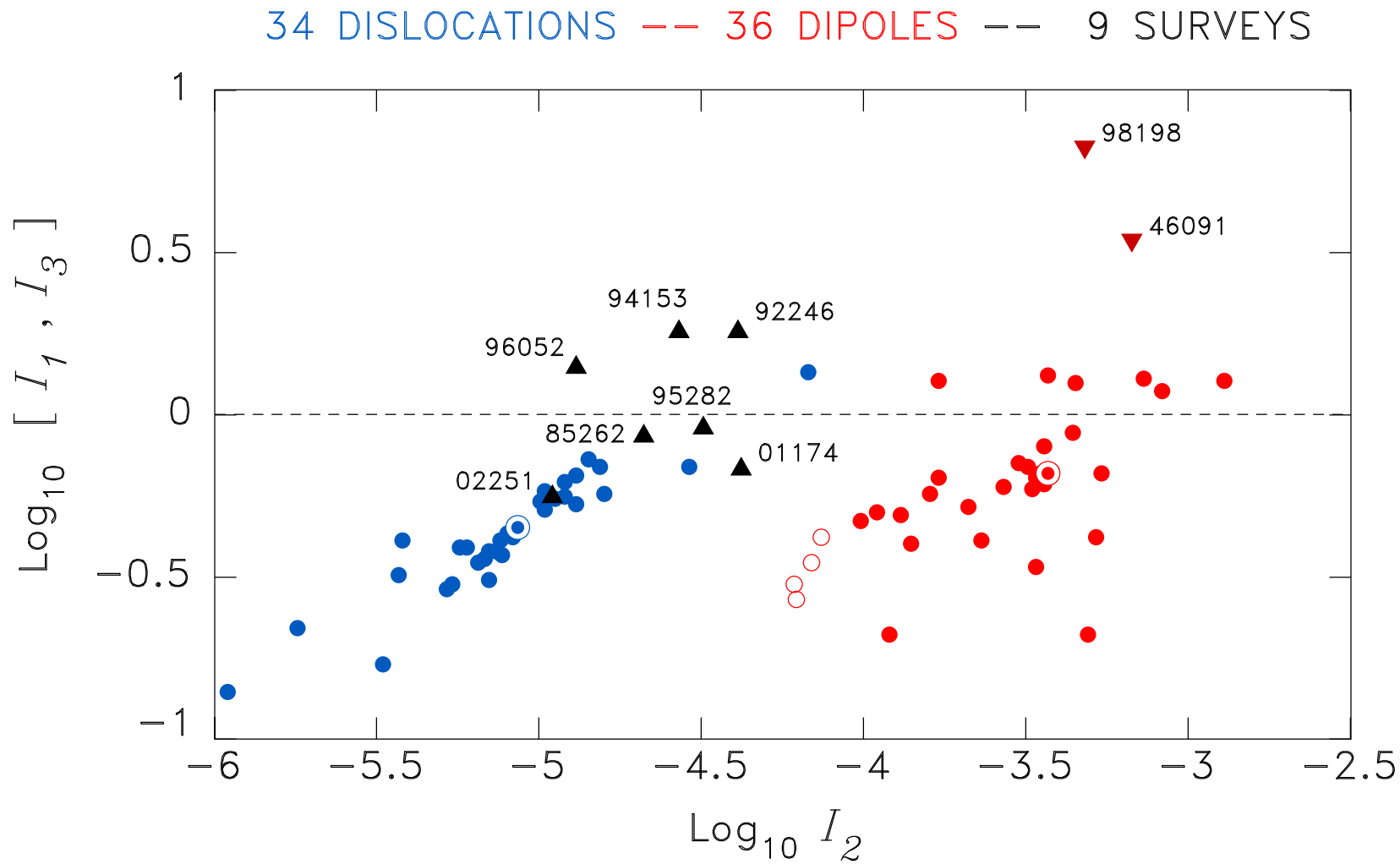
ASPECT RATIO OF RUN-UP DISTRIBUTION ALONG BEACH

[Okal and Synolakis, in press]





MAX. RUN-UP SCALED TO FAULT SLIP
MAX. RUN-UP SCALED TO INITIAL TROUGH



ASPECT RATIO OF RUN-UP DISTRIBUTION ALONG BEACH

[Okal and Synolakis, in press]

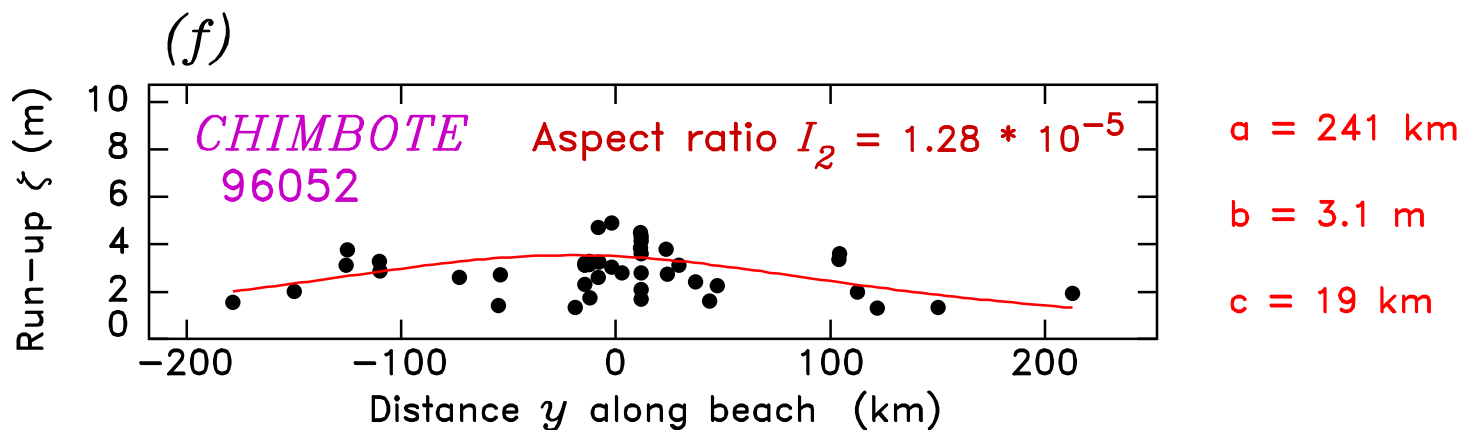
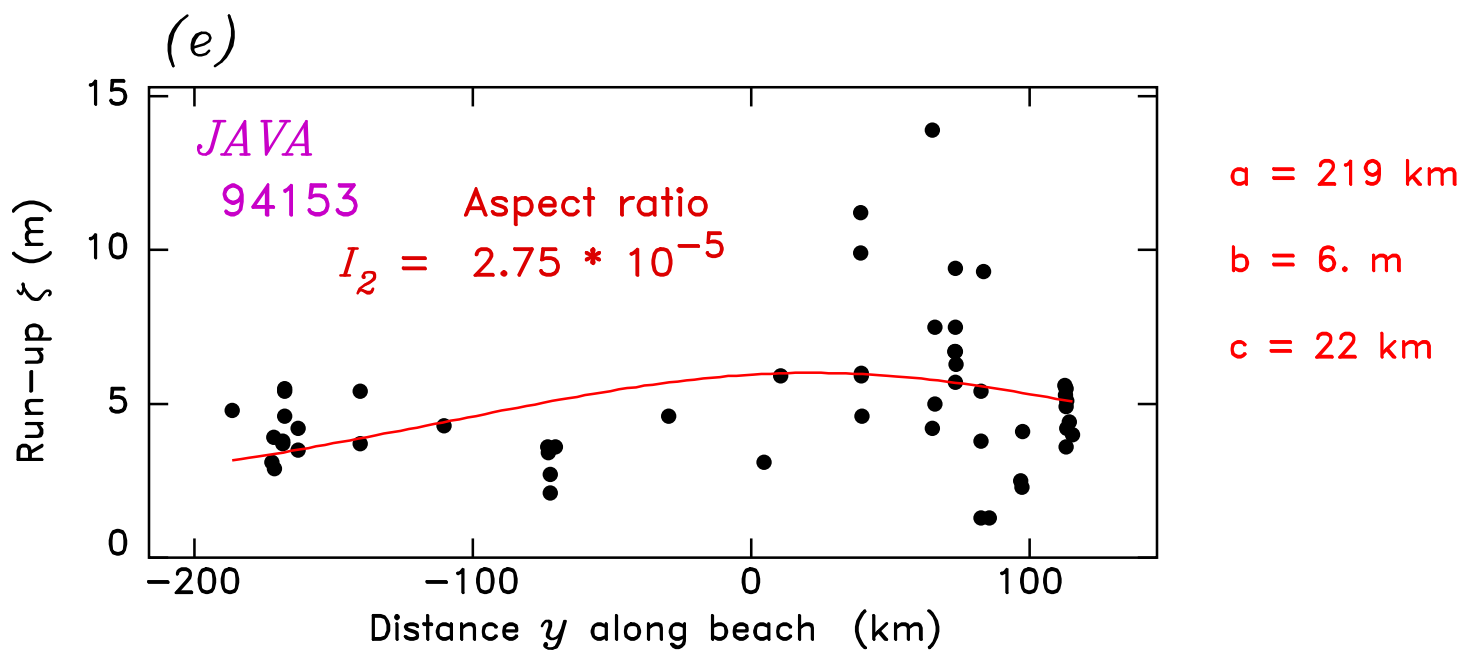
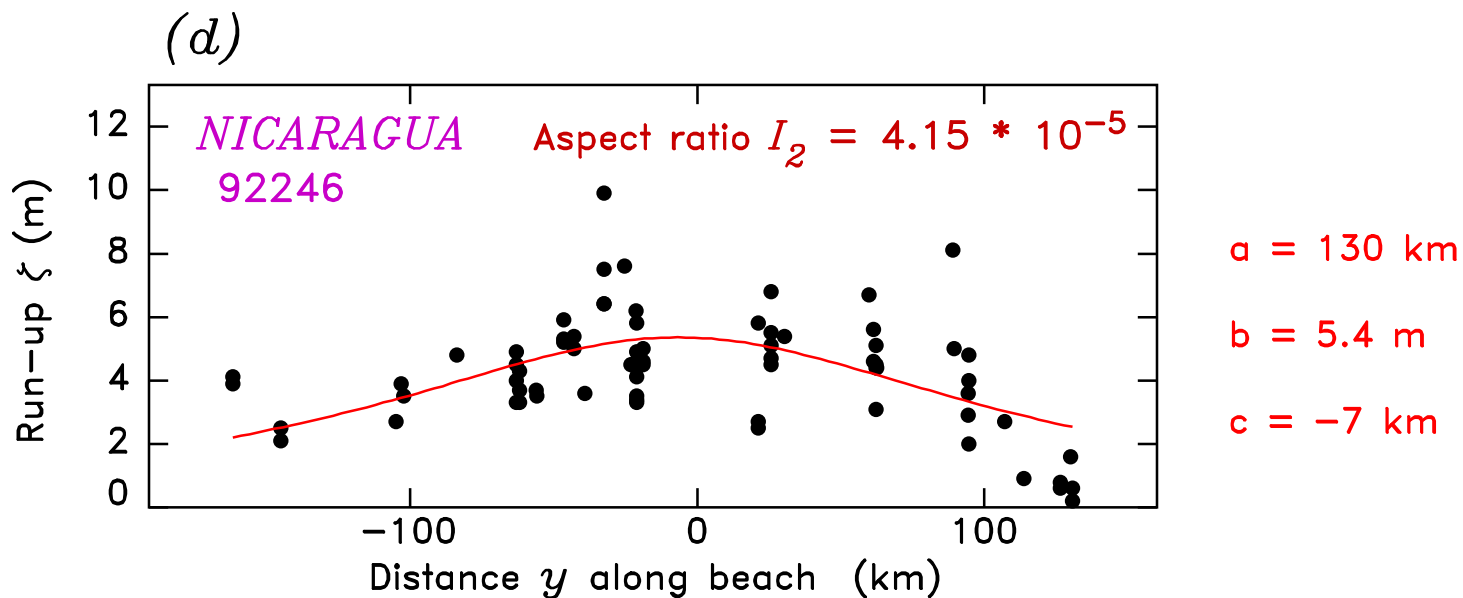


FIGURE 6 d-f

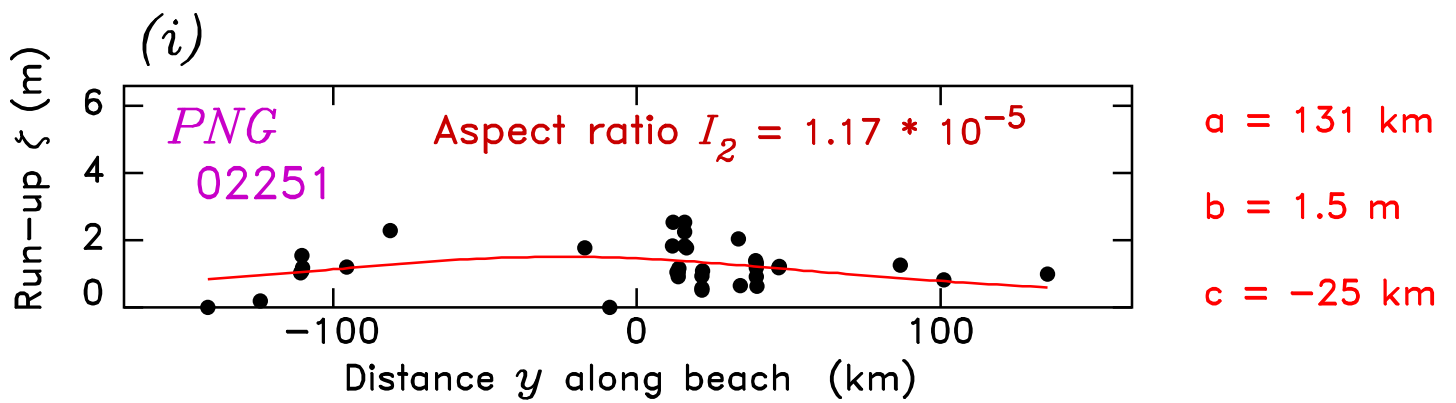
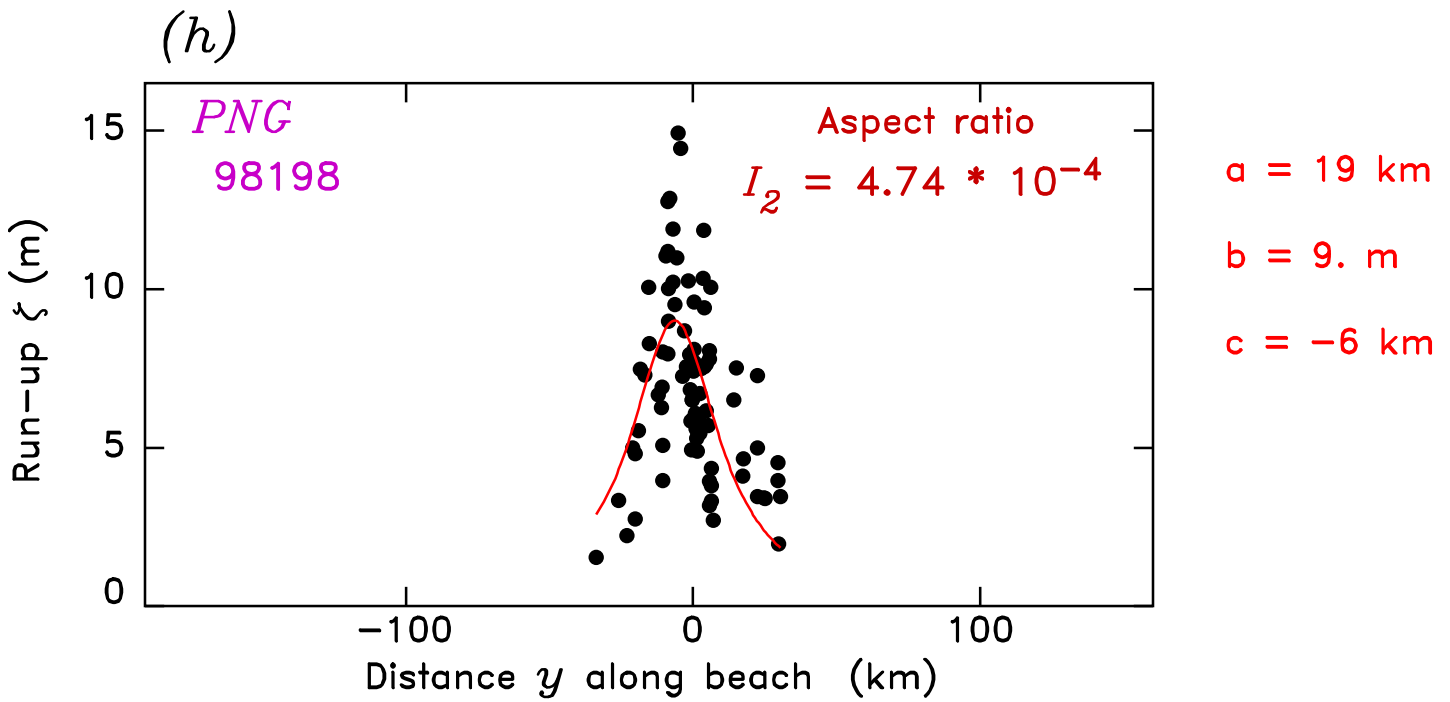
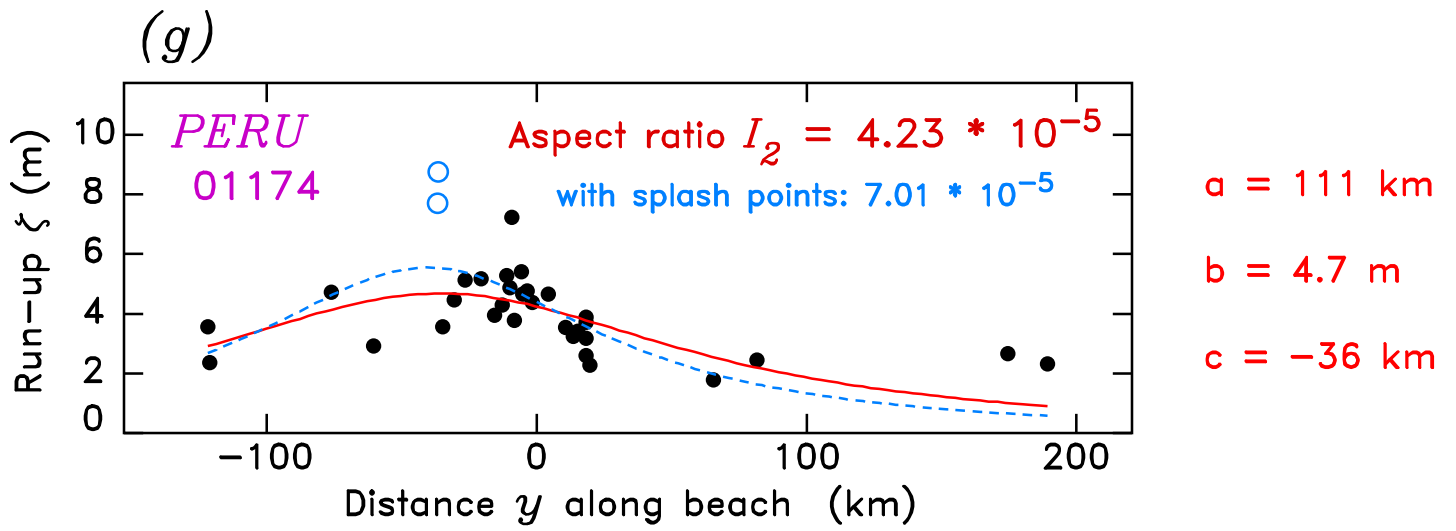
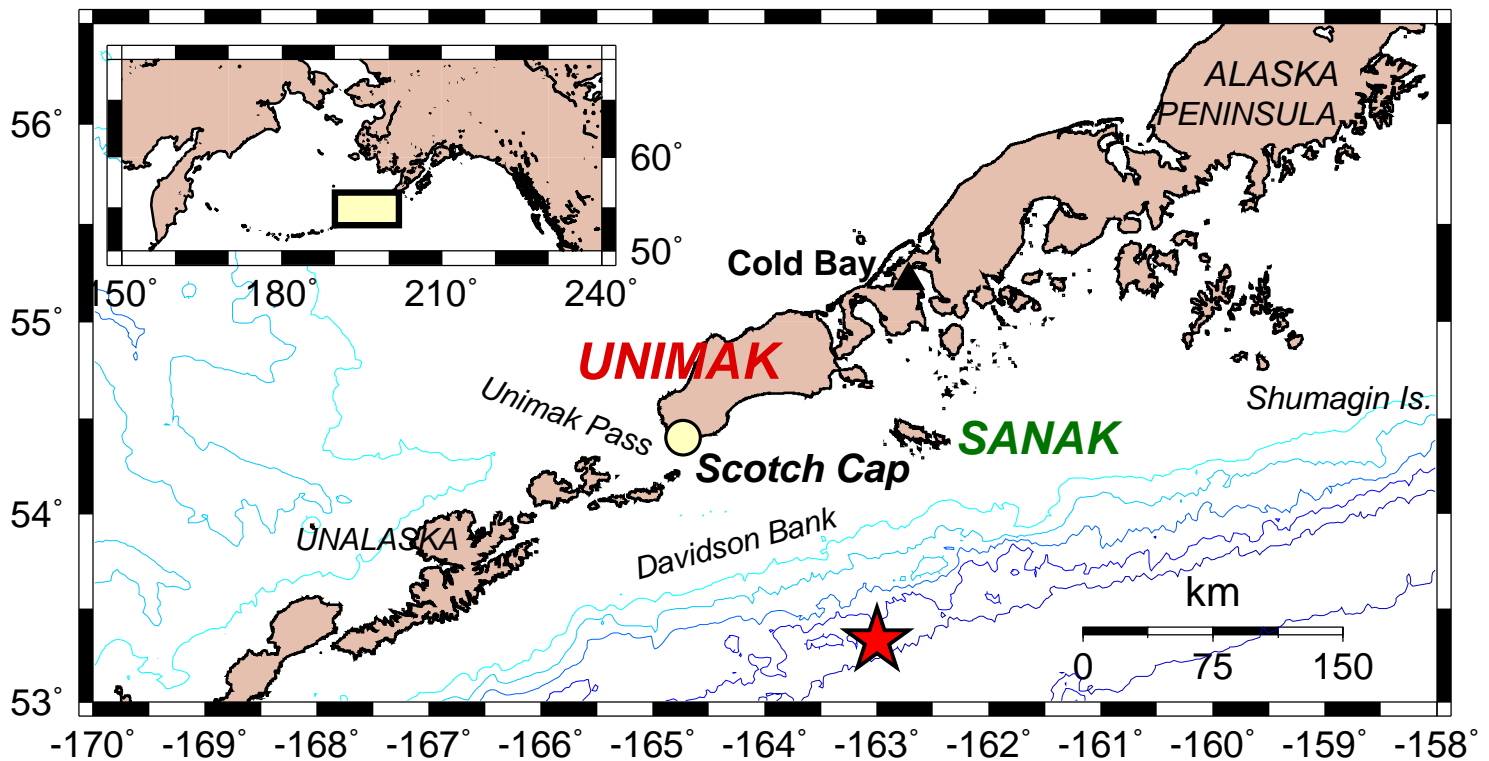


FIGURE 6g-i

THE 1946 ALEUTIAN TSUNAMI: A PERSISTING CHALLENGE

- A rather moderate earthquake ($M_{PAS} = 7.4$)
- A devastating transpacific tsunami
- A catastrophic local tsunami

Scotch Cap lighthouse eradicated.



THE QUESTION REMAINS

How to model the source of the tsunami: A gigantic earthquake source, or a large underwater landslide, triggered by the seismic event?

DESTRUCTION OF THE LIGHTHOUSE AT SCOTCH CAP, UNIMAK Is.

[Photog. *H. Hartman*; Courtesy G. Fryer]

Before (1945)



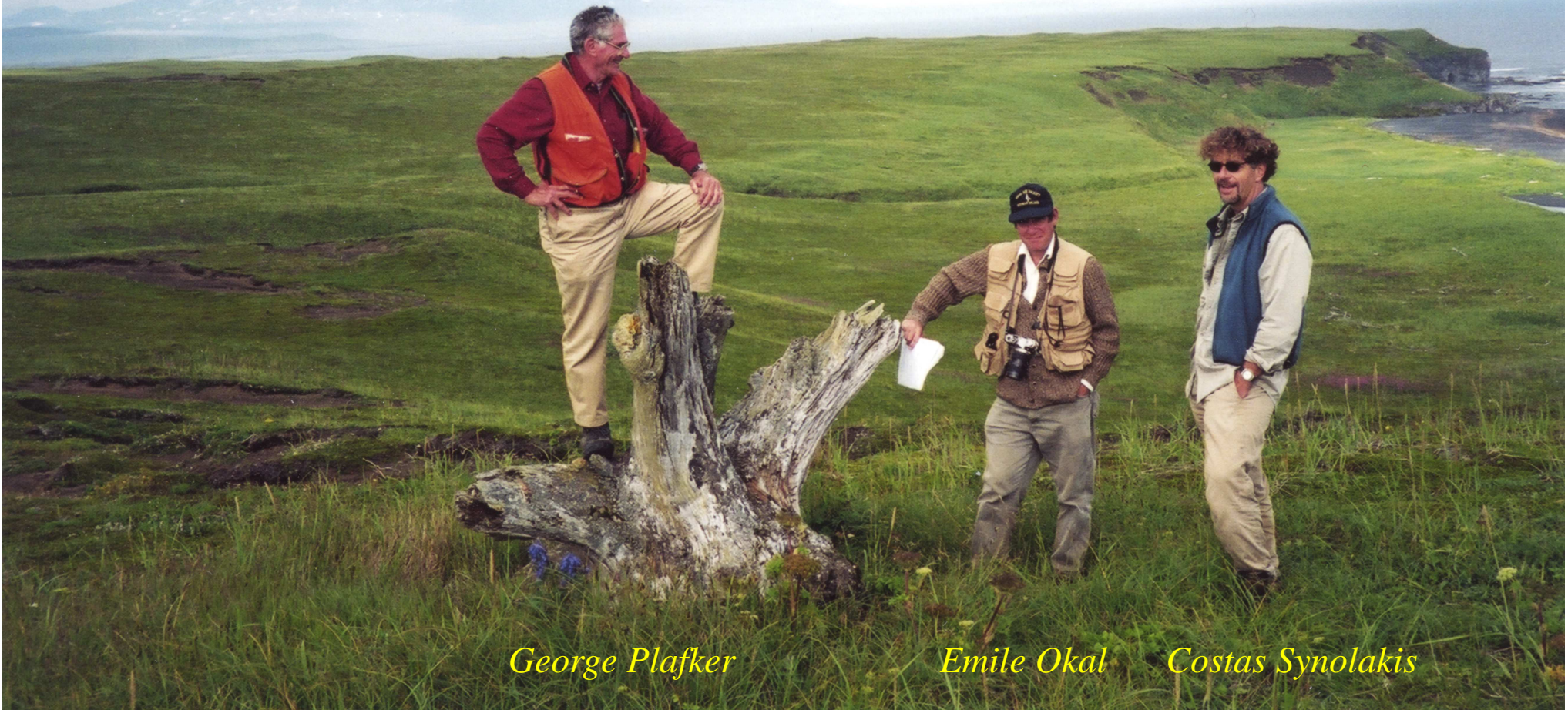
After (est. 03-04 (?) Apr. 1946)



No trees grow on the Eastern Aleutian Islands...

Thus, large logs lying several hundred meters inland at altitudes of 10 to 30 m constitute watermarks of inundation by a tsunami, since they are way beyond the limit of even the most powerful storm surges.

In recent decades, only the 1946 tsunami is a viable candidate as the agent of their deposition.



George Plafker

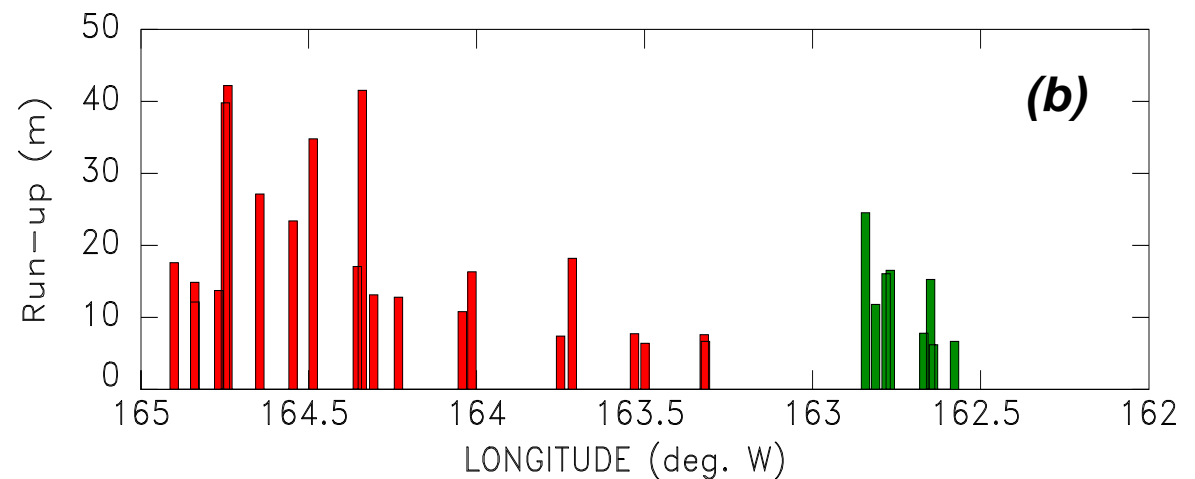
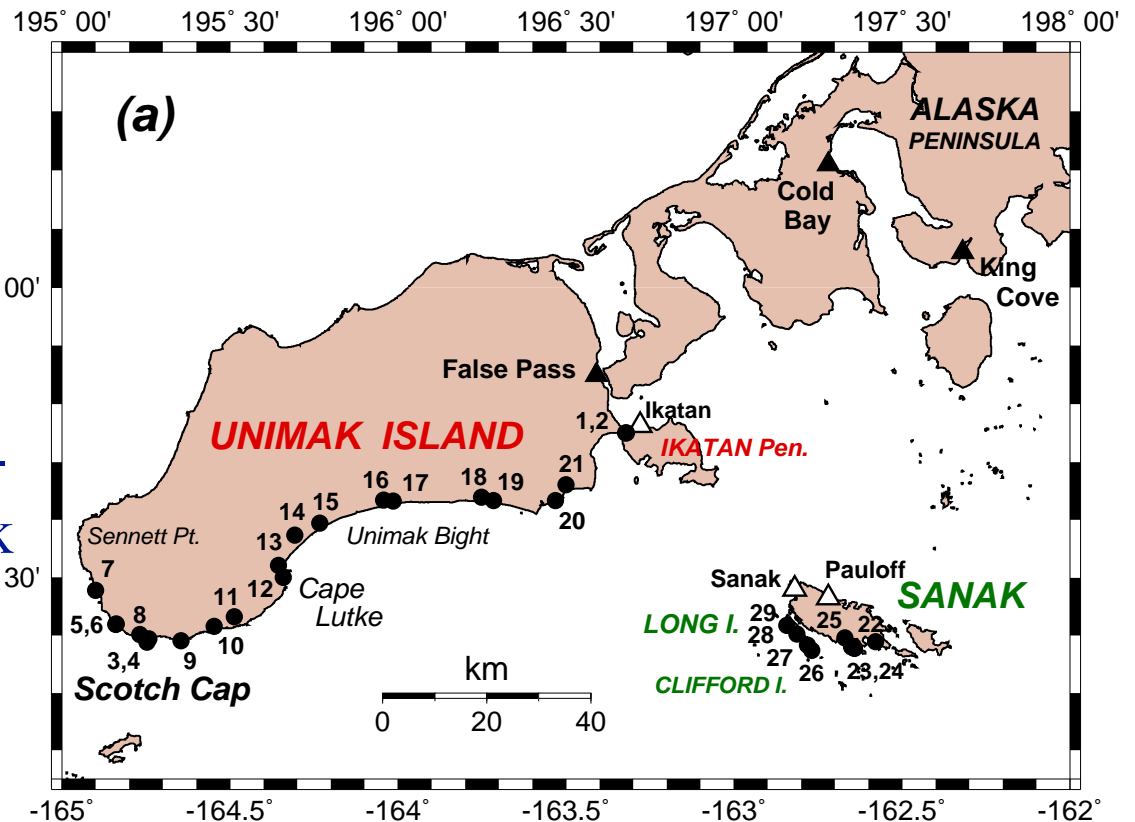
Emile Okal

Costas Synolakis

Cape Lutke, UNIMAK ISLAND

1946 RESULTS IN NEAR FIELD

- Run-up at Scotch Cap: 42 m (Ruins of Radio Station)
- Extreme run-up concentrated along 40 km of coast line.
- Run-up "only" 15 m, but inundation up to 2 km along Unimak Bight
- Run-up up to 24 m on Sanak



Near-field *Aspect Ratio* of Run-up Distribution at Unimak (6.4×10^{-4}) even larger than for PNG-1998, thus

REQUIRING LANDSLIDE SOURCE

