Benchmark problem 3 The third international workshop on long-wave runup models Catalina Island, June 2004

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Theory employed

The models

- Eulerian FDM for standard Boussinesq equations Nonlinear code without particular run-up feature Used only in linear mode, with and without dispersion terms
- Lagrangian FDM for Boussinesq type equations Method particularly designed for runup Fully nonlinear, but weakly dispersive. Employed in hydrostatic and dispersive mode

Definitions

- In equilibrium depth, $h' = \partial h(x,t) / \partial x$ etc.
- η surface elevation
- $H = h + \eta$ total depth
- \overline{u} depth-averaged horizontal velocity

Averaged material derivate

$$rac{\mathrm{D}}{\mathrm{D}t} = rac{\partial}{\partial t} + \overline{u}rac{\partial}{\partial x}$$

The fully nonlinear Boussinesq equations

$$\begin{split} &\frac{\mathrm{D}H}{\mathrm{D}t} = -H\frac{\partial\overline{u}}{\partial x}, \\ &\left(1 - \frac{1}{2}Hh'' - h'\frac{\partial\eta}{\partial x}\right)\frac{\mathrm{D}\overline{u}}{\mathrm{D}t} = \\ &-g\frac{\partial\eta}{\partial x} - \frac{1}{3}H\frac{\partial}{\partial x}\left(\frac{\mathrm{D}^{2}H}{\mathrm{D}t^{2}}\right) - \frac{2}{3}\frac{\partial H}{\partial x}\frac{\mathrm{D}^{2}H}{\mathrm{D}t^{2}} \\ &+ \frac{h'}{H}\left(\frac{\mathrm{D}H}{\mathrm{D}t}\right)^{2} - h''\overline{u}\frac{\mathrm{D}H}{\mathrm{D}t} + \left(\frac{\partial\eta}{\partial x}h'' + \frac{1}{2}Hh'''\right)u^{2} + S_{1} + S_{2} \end{split}$$

Red: "Important" dispersion terms Green: fully nonlinear version, $S_1 + S_2$: source terms.

Long waves and Lagrangian coordinates

Simple flow structure \Rightarrow material fluid columns remain (nearly) vertical Lagrangian coordinate a: $\frac{Da}{Dt} = 0$, a(x, 0) = xTransformation $\frac{D}{Dt} \rightarrow \frac{\partial}{\partial t}$, $\frac{\partial}{\partial x} \rightarrow \frac{H}{H_0} \frac{\partial}{\partial a}$ etc., $H_0 = H(a, 0)$

$$egin{aligned} &Hrac{\partial x}{\partial a}=H_{0},\ &(1-()-())rac{\partial \overline{u}}{\partial t}=-grac{H}{H_{0}}rac{\partial H}{\partial a}+grac{\partial h}{\partial x}-()+()+S_{1}+S_{2},\ &rac{\partial x}{\partial t}=\overline{u}, \end{aligned}$$

Shoreline: H = 0 at fixed a

The slide representation

Momentum conservation

$$(1-()-())rac{\partial\overline{u}}{\partial t}=-grac{H}{H_0}rac{\partial H}{\partial a}+grac{\partial h}{\partial x}-()+()+rac{S_1}{S_1}+S_2,$$

dominant explicit forcing term: $g \frac{\partial h}{\partial x}$

Additional h.o. (dispersive) terms from slide

$$egin{aligned} S_1+S_2&=rac{1}{2}Hrac{\partial^3 h}{\partial t^2\partial x}-rac{1}{g}\left(2\overline{u}rac{\partial^2 h}{\partial t\partial x}+rac{\partial^2 h}{\partial t^2}
ight)rac{\partial\overline{u}}{\partial t}\ &-rac{\partial^2 h}{\partial t\partial x}rac{\partial H}{\partial t}+H\overline{u}rac{\partial^3 h}{\partial t\partial x^2} \end{aligned}$$

The Lagrangian Boussinesq model

- Second order FDM model. Staggered grid in space and time.
- *H*-node at shoreline position (fixed *a*). Condition $H \equiv h + \eta = 0$ implemented directly.
- Hydrostatic (NLSW) version is explicit.
- Dispersive version implicit iteration on nonlinearity (may be avoided)
- Wave paddle is also simple to implement.
- No smoothing or filtering for non-breaking waves.
- Hydrostatic (NLSW) extensions to non-planar waves exist

Results

Nomenclature

- ref.: (Semi-)analytic solutions from reference papers
- Bouss(NL) : Nonlinear Boussinesq solution (Lagrangian)
- NLSW: Nonlinear shallow water solution (Lagrangian)
- Bouss(L): Linear Boussinesq solution (Eulerian)
- SW(L): Linear shallow water solution (Eulerian)
- pot(NL): Full potential theory (boundary integral method)

Color codes are preferably kept consistent.

Case B, t = 4.5



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Case A, t = 0.1



A at t = 0.1. Comparison of models.

Case A, t = 1.5



A at t = 1.5. Comparison of models.

Case B, t = 1.0



B at t = 1.0. Comparison of models.

Case B, t = 2.5



B at t = 2.5. Comparison of models.

Case B, t = 4.5



B at t = 4.5. Comparison of models.

A, t = 0.5: convergence



Bouss(NL), A: t = 0.5. Curves marked by Δa (m)

B, t = 2.5: convergence



Bouss(NL), B: t = 2.5. Curves marked by Δa (m)

Concluding remarks 1

Wave models

- Linear h.o. term for slide (S_1) affects wave generation slightly for case B
- Dispersion important for B when generated waves propagate into deeper water
- Even case A is affected by nonlinearity close to beach
- Linear models perform fairly well also for case B, but significant errors (save at wave front for small t).
- case A: Reference (analytic solution) superior to "naive" SW close to shore. (Not the case for B)

Concluding remarks 2

Performance of techniques

(Concerning the computed time span)

- Shoreline region most sensitive to resolution
- Case A most demanding due to large difference between length scales at beach and in deeper water.
- Less demanding than benchmark 1
- Problems in store for case A and t > 1.5?

Extras

The simulations

- Series of resolutions employed for each model
- Eulerian FDM: adaptive spatial refinement employed
- Lagrangian FDM: uniform resolution initially
- Different approximations in source term tested

Case A, t = 0.5



A at t = 0.5. Comparison of models.

Case A, t = 1.0



A at t = 1.0. Comparison of models.

Case B, t = 0.5



B at t = 0.5. Comparison of models.

A, t = 1.5: convergence



Bouss(NL), A: t = 1.5. Curves marked by Δa (m)

B, t = 0.5: convergence



Bouss(NL), B: t = 0.5. Curves marked by Δa (m)

h.o. source terms



Bouss(NL), B: t = 4.5. Source terms

h.o. source terms



Bouss(NL), A: t = 1.5. Source terms

References on models and related papers

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